



Stability analysis of numerical interface conditions in coupled ocean-atmosphere models

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Overview

Motivation

Current coupling algorithms for climate models are equivalent to performing a single step of an iterative procedure [Lemarie et al. 2015]. This may not be enough to ensure numerical stability.

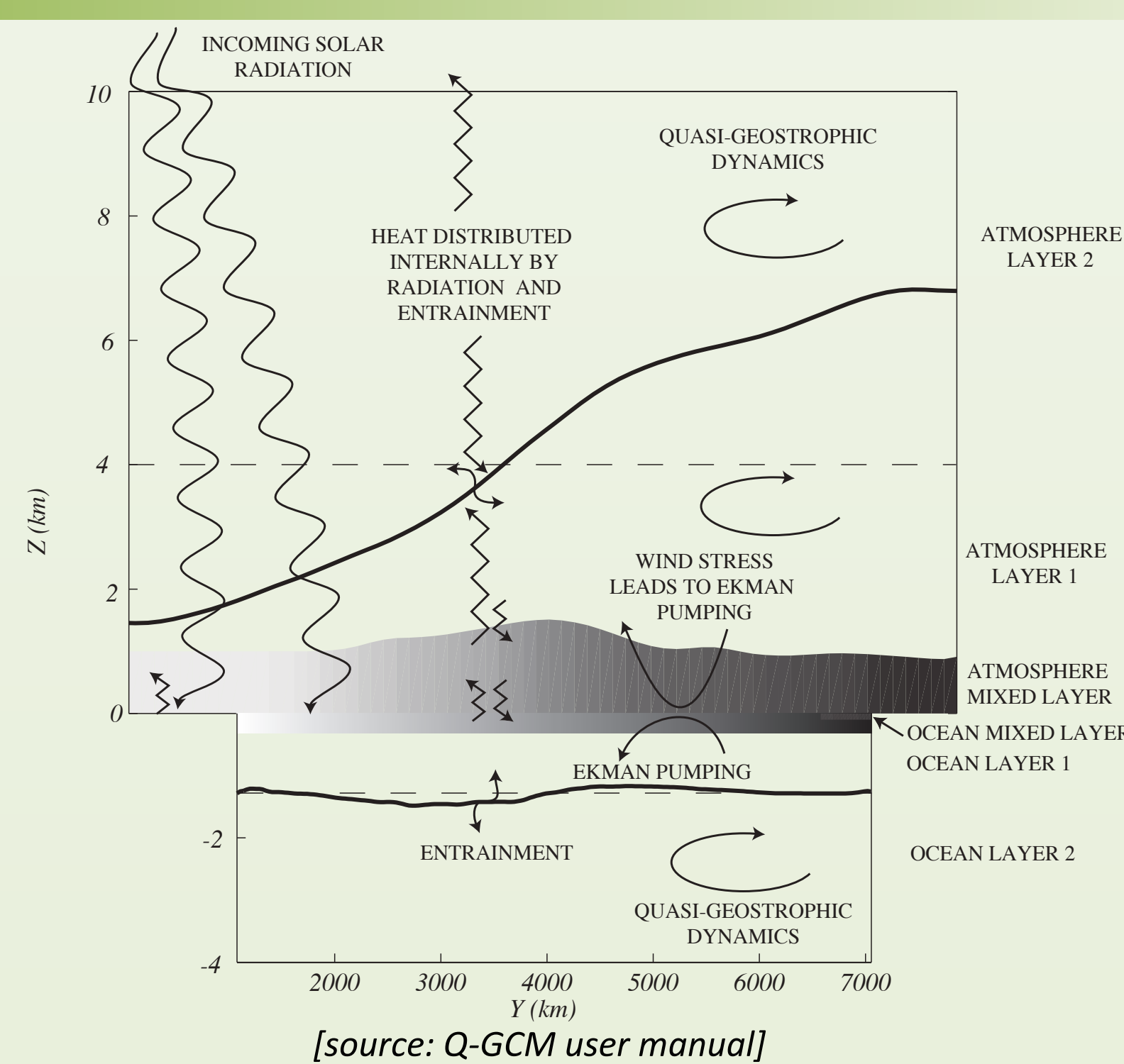
Objectives

- Perform rigorous stability analysis of interface conditions for a variety of coupling algorithms
- Gain insight on how to improve the existing methods or propose new coupling schemes

Model Exemplars: Q-GCM

The Quasi-Geostrophic Coupled Model

- Intermediate-complexity mid-latitude coupled climate model
- Candidate testbed for comparing different strategies and analyzing instability



Coupling mechanism

- Includes mixed layers to allow the exchange of heat flux and momentum
- At the interface, **Ekman pumping** drives momentum from atmosphere to ocean; latitudinal variations in solar radiation drives surface **heat flux** in both direction

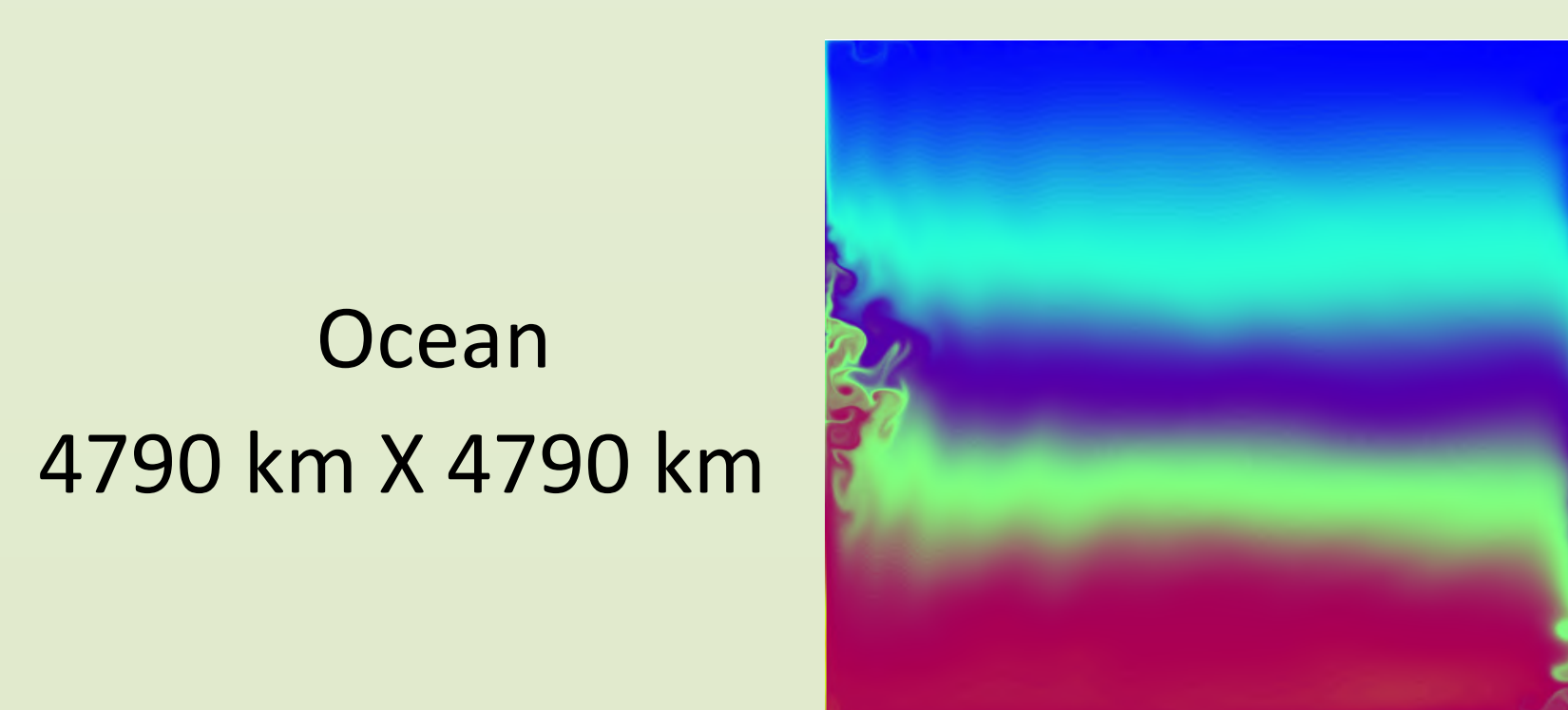
$${}^a q_{kt} + ({}^a u_k {}^a q_k)_x + ({}^a v_k {}^a q_k)_y = \frac{f_0 ({}^a e_k - {}^a e_{k-1})}{{}^a H_k} - \frac{{}^a A_4 \nabla_H^2 {}^a p_k}{f_0}$$

$${}^o q_{kt} + ({}^o u_k {}^o q_k)_x + ({}^o v_k {}^o q_k)_y = \frac{f_0 ({}^o e_k - 1 - {}^a e_k)}{{}^o H_k} + \frac{{}^o A_2 \nabla_H^2 {}^o p_k}{f_0} - \frac{{}^o A_4 \nabla_H^2 {}^o p_k}{f_0}$$

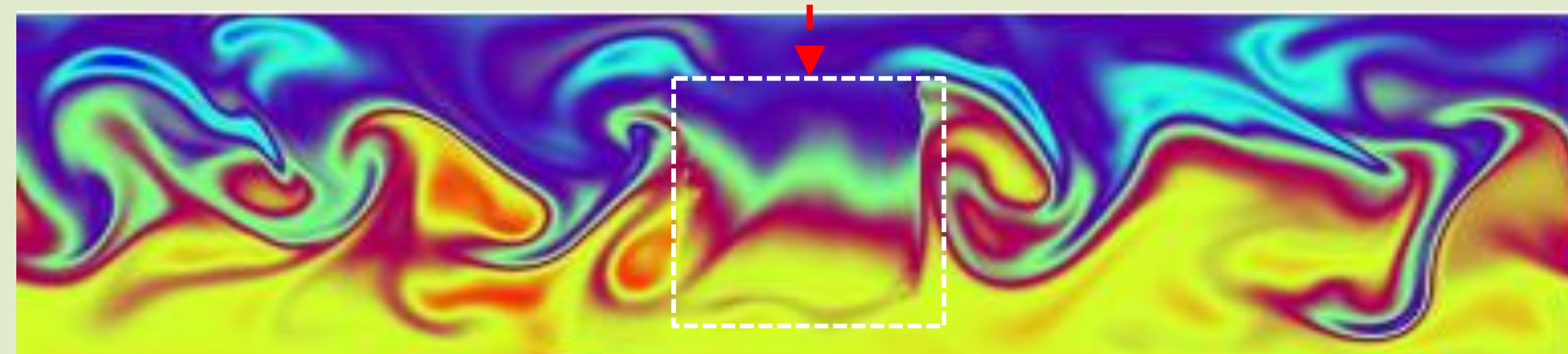
notation: k - layer number a - atmosphere o - ocean
q - vorticity p - pressure e - entrainment

$${}^a T_{mt} + ({}^a u_m {}^a T_m)_x + ({}^a v_m {}^a T_m)_y + \frac{{}^a w_{ek} {}^o T_m}{{}^a H_m} = {}^a K_2 \nabla_H^2 {}^a T_m - {}^a K_4 \nabla_H^4 {}^a T_m + \frac{-{}^a F_m + {}^o F_0}{\rho {}^a C_p {}^a h_m}$$

$${}^o T_{mt} + ({}^o u_m {}^o T_m)_x + ({}^o v_m {}^o T_m)_y + \frac{{}^o w_{ek} {}^o T_m}{{}^o H_m} = {}^o K_2 \nabla_H^2 {}^o T_m - {}^o K_4 \nabla_H^4 {}^o T_m + \frac{-{}^o F_m + {}^a F_0}{\rho {}^o C_p {}^o h_m}$$



Atmosphere
30640 km X 7600 km



Surface Temperature for the double gyre example (6 months)

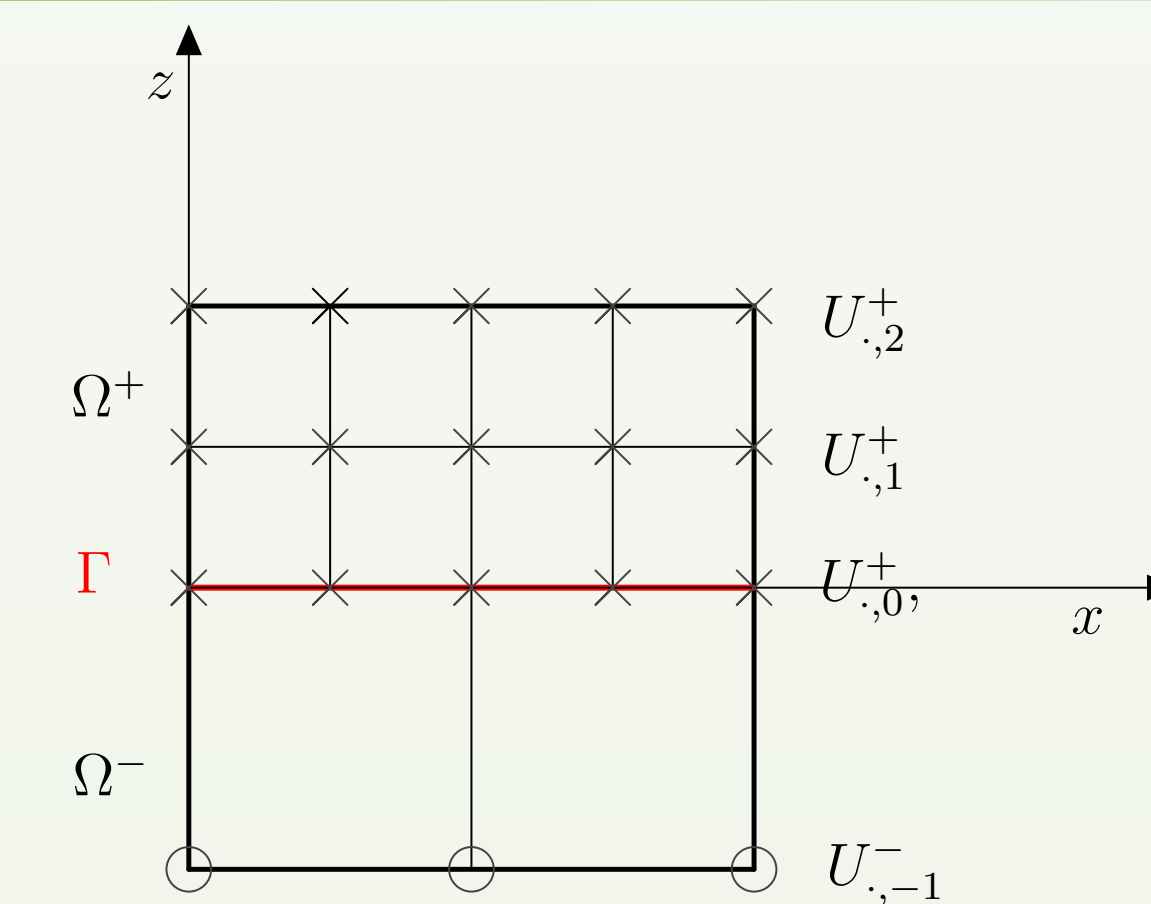
Observations

- Running in the coupled mode, the ocean model stays stable with a timestep size of up to 40 minutes; running in the ocean-only mode, the timestep size can be larger (>72 minutes)

Interface Conditions

Practical assumptions

- Both components can be run simultaneously to promote task parallelism
- Ocean step size is a multiple of atmosphere step



Dirichlet-Neumann condition

$$\begin{cases} \mathcal{L}^+ U^+ = f^+ & \text{in } \Omega^+ \times [0, \mathcal{T}] \\ U_0^+ = U_0^- & \text{on } \Gamma \times [0, \mathcal{T}] \\ \mathcal{L}^- U^- = f^- & \text{in } \Omega^- \times [0, \mathcal{T}] \\ \mathcal{F}(U_0^-) = \mathcal{F}(U_0^+) & \text{on } \Gamma \times [0, \mathcal{T}] \end{cases}$$

Bulk interface condition

$$\begin{cases} \mathcal{L}^+ U^+ = f^+ & \text{in } \Omega^+ \times [0, \mathcal{T}] \\ \rho^+ \mathcal{F}(U^+) = \tau & \text{on } \Gamma \times [0, \mathcal{T}] \\ \mathcal{L}^- U^- = f^- & \text{in } \Omega^- \times [0, \mathcal{T}] \\ \rho^- \mathcal{F}(U^-) = \tau & \text{on } \Gamma \times [0, \mathcal{T}] \end{cases}$$

Discretize for 1D Diffusion equation

$$\begin{aligned} U_{j,j}^{+,n+1} - U_{j,j}^{+,n} &= \left(\frac{\nu_+ \Delta t}{(\Delta z_+)^2} \right) (U_{j,j+1}^{+,n} - 2U_{j,j}^{+,n} + U_{j,j-1}^{+,n}), & j < 0 \\ U_{j,j}^{+,n+1} - U_{j,j}^{+,n} &= \left(\frac{\nu_+ \Delta t}{(\Delta z_+)^2} \right) (U_{j,j+1}^{+,n} - 2U_{j,j}^{+,n} + U_{j,j-1}^{+,n}) + \frac{\nu_+ \Delta t}{(\Delta z_+)^2} (U_{j,j+1}^{+,n} - U_{j,j-1}^{+,n}), & j \geq 1 \\ U_{j,0}^{-,n+1} - U_{j,0}^{-,n} &= \Delta t \left(\frac{\mathcal{F}_{j,0}^{+,n} - \mathcal{F}_{j,0}^{-,n}}{\Delta z_+ + \Delta z_-} \right), & \text{where } \mathcal{F} \text{ is the flux} \\ U_{j,j}^{-,n+1} - U_{j,j}^{-,n} &= \left(\frac{\nu_- \Delta t}{(\Delta z_-)^2} \right) (U_{j,j+1}^{-,n} - 2U_{j,j}^{-,n} + U_{j,j-1}^{-,n}), & j \leq -2 \\ U_{j,j}^{-,n+1} - U_{j,j}^{-,n} &= \left(\frac{\nu_- \Delta t}{(\Delta z_-)^2} \right) (U_{j,j+1}^{-,n} - 2U_{j,j}^{-,n} + U_{j,j-1}^{-,n}) + \frac{\nu_- \Delta t}{\Delta z_-} \left(\frac{\rho_2}{\rho_1} \alpha (U_{j,j}^{+,n} - U_{j,j}^{-,n}) - \frac{\nu_-}{\Delta z_-} (U_{j,j}^{-,n} - U_{j,j-1}^{-,n}) \right), & j = -1 \end{aligned}$$

Normal Mode Analysis

For explicit Euler consider the solution

$$U_j^n = \begin{cases} \mathcal{A}^n k_-^j, & j = 0, -1, -2, \dots \\ \mathcal{A}^n k_+^j, & j = 1, 2, \dots \end{cases}$$

$$\begin{aligned} A &= 1 + d_+ (k_+ - 2 + k_+^{-1}) \\ A &= 1 + 2d_+ r_+ (k_+ - 1) - 2d_- r_- (1 - k_-^{-1}) \\ A &= 1 + d_- (k_- - 2 + k_-^{-1}) \end{aligned}$$

Define $r_{\pm} = \frac{\Delta z_{\pm}}{\Delta z_+ + \Delta z_-}$ $d_{\pm} = \frac{\nu_{\pm} \Delta t}{(\Delta z_{\pm})^2}$

Solve for k and choose negative root $1 = r_+ \left(1 - \sqrt{1 + \frac{4d_+}{A-1}} \right) + r_- \left(1 - \sqrt{1 + \frac{4d_-}{A-1}} \right)$

Asymptotic solution of the requirement $|A| < 1$ yields

- If $\Delta z_- \ll \Delta z_+$ then $d_+ < \frac{1}{2}$
- If $\Delta z_+ \ll \Delta z_-$ then $d_- < \frac{1}{2}$
- If $d_{\pm} \gg 1$ then $\frac{\nu_-}{\nu_+} \approx 1$

For backward Euler we have to use $U_{j,0^+}^{+,n+1} = U_{j,0^-}^{-,n}$

Thus consider the solution

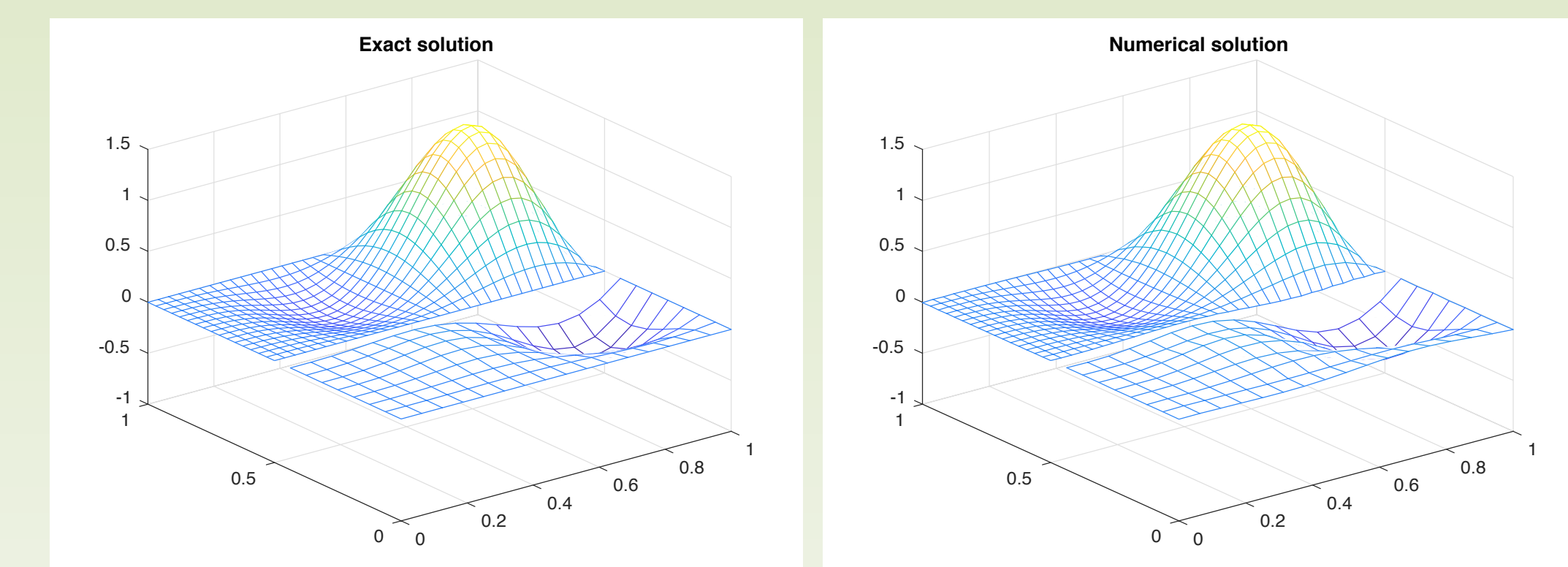
$$U_j^n = \begin{cases} \mathcal{A}^n k_-^j, & j = 0^-, -1, -2, \dots \\ \mathcal{A}^{n-1} k_+^j, & j = 0^+, 1, 2, \dots \end{cases}$$

- If $\Delta z_- \ll \Delta z_+$ the scheme is stable
- Otherwise the scheme can be unstable

Analysis on bulk interface conditions also implies additional constraints on step size

Conclusions and Future Work

- Instability originates from interface condition and the implementation
- Validated on an advection-diffusion model
- Extending to a passive tracer model
- Adding the iterative Schwarz method in Q-GCM to improve the step size



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