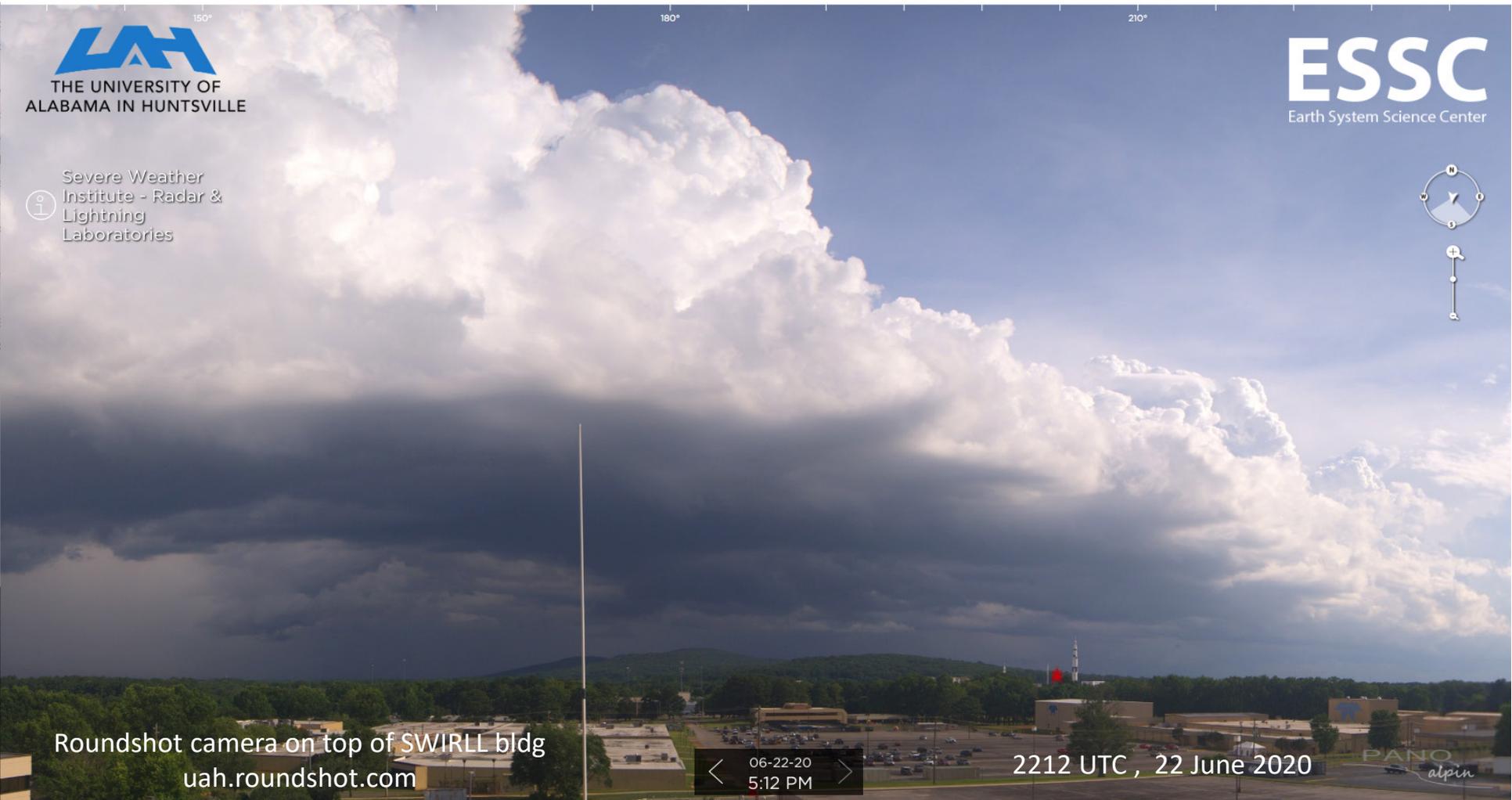


# Boundary layer heterogeneity and deep convection

(relevance to AMF3 SEUS)



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# Background statement

- This presentation summarizes methodologies and related science objectives in current research designed to understand the interaction between deep convection and spatial/temporal variability within the PBL (BL). The domain is our back yard, northern Alabama.
- VORTEX-SE campaigns, 2016-current

# Overview

## 1. Network design challenges

- Define the scientific objectives, then . . .
- arrange radar/profiler resources to best satisfy the scientific requirements
- within the constraints produced by obstacles (trees, orography) which limit the number of sites

## 2. Science: Understand BL heterogeneity and evolution

- How it impacts Convective Initiation and QLCS evolution
- How it modulates deep convection

*General goal: Obtain both storm-scale and BL flow over (part of) the same mesoscale domain*

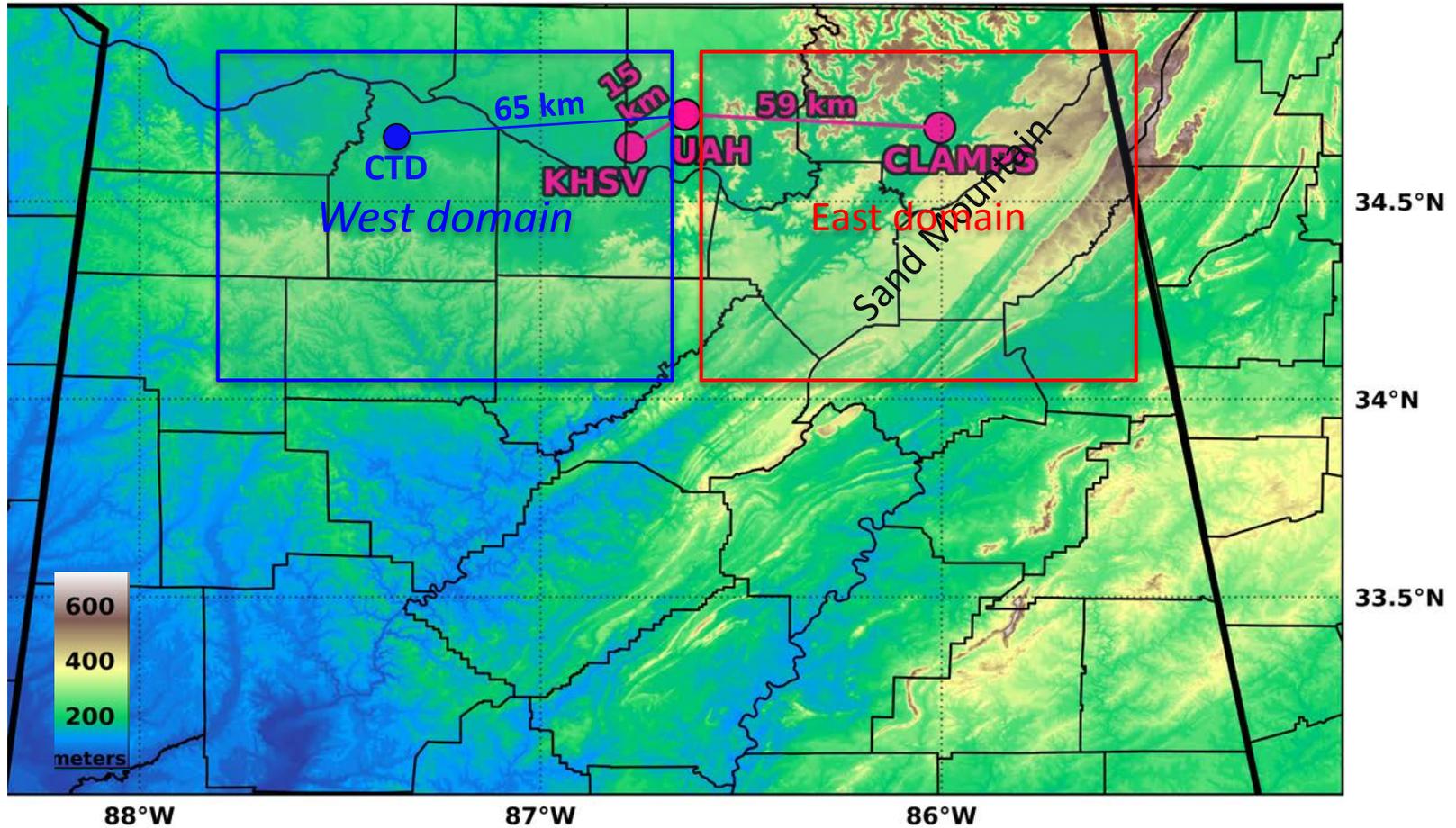
# Background: Seasonal variation of SE deep convection

- Core warm season (Jun-Aug)
  - Diurnal dominance, afternoon maximum
  - Low shear, high CAPE (Convective Available Potential Energy)
  - Transient convective cells, occasional upscale growth to MCS (refer to title slide)
- Core cold season (Dec-Feb)
  - Strong synoptic-scale forcing (cyclone passages)
  - Organized convection within cyclone warm sector: Quasi-linear Convective System (QLCS, long-lived) is the dominant convective mode
    - High shear low CAPE
  - (Non-convective) stratiform precipitation in regions of greater stability (e.g., north of the cyclone)
- Spring and Autumn seasons (Mar-May, Sep-Nov)
  - Mix of warm/cold season, but QLCS will prevail
  - Tropical cyclones occasionally affect the SE (June-November)

# 1. Deployment challenges

- Trees and orography present deployment challenges for any system requiring good view of the horizon.
  - Notable examples:
    - Scanning remote sensing platforms: radars, lidars
    - Photography
    - Balloon soundings – not as restrictive
  - Considerable time may be invested in “scouting out” potential locations
  - Any radar site that has limited blockage below about 1.5 deg elevation angle is considered very good for Northern AL
  - Excellent example of a good site: Courtland AP (3 slides down)
- Road network – applies to mobile operations and design of adaptable networks of profiling systems for a specific IOP (science objective)

# Northern Alabama domain



- Long-term profiling sites for W domain: UAH and Courtland (CTD, 2019-2020)
- Science focus on BL heterogeneity produced by variations in surface fluxes, surface roughness length, low-level wind/thermo variations
- Long-term profiling sites for E domain: UAH and CLAMPS (2017)
- Science focus: topographic influence on BL flows
- Challenging: complex flows, subpar road network, limited radar sites



# Courtland Airport (CTD)

Instruments during 2019-2020 campaign

**Legend**

- 449/RASS
- CRN
- DWL
- MAX
- PA0
- Sfc/GPS
- Twr



Profiling systems

Lockheed-Martin (DOD)

MAX radar



# Ground-level views from the CTD site

2255 UTC

CI over & just  
NE of CTD  
2254 UTC,  
4/23/20



View to NE, trees in foreground

2325 UTC

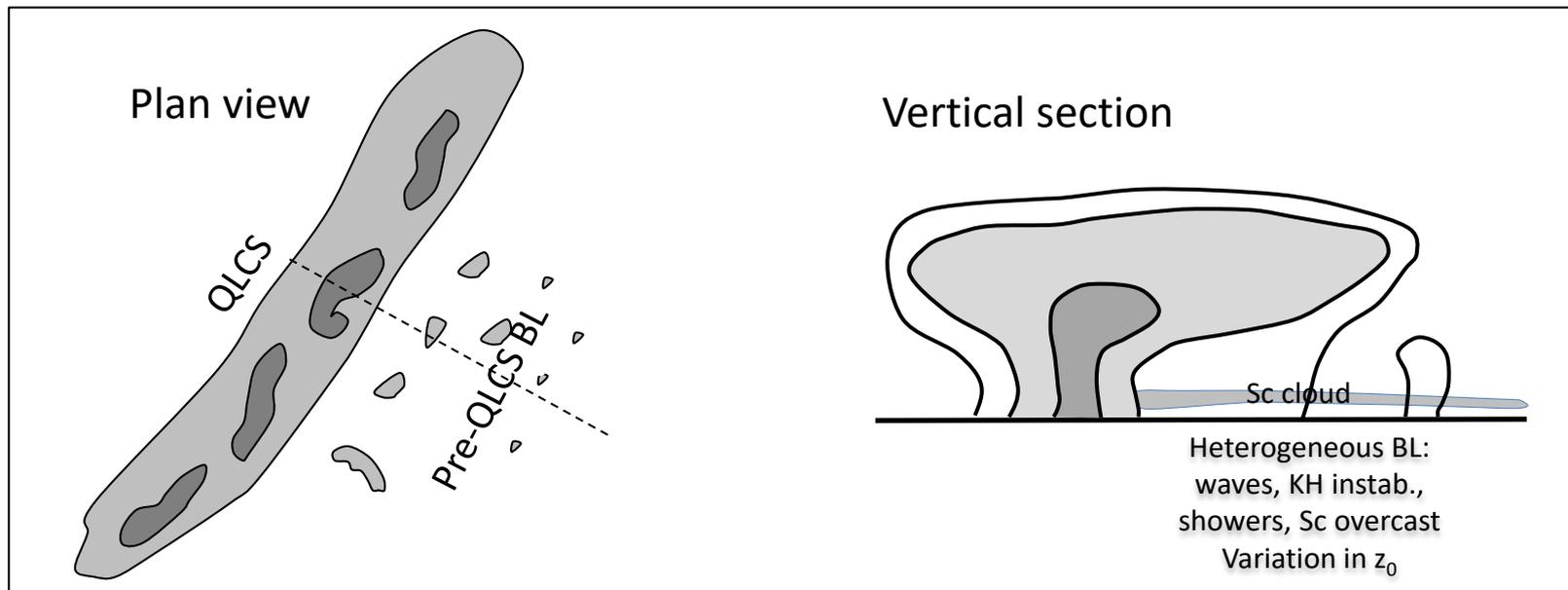


View to SE, open airport area in foreground

Newly-formed, modest  
QLCS receding from CTD  
2325 UTC, 4/23/20

## 2. Science objectives (current research)

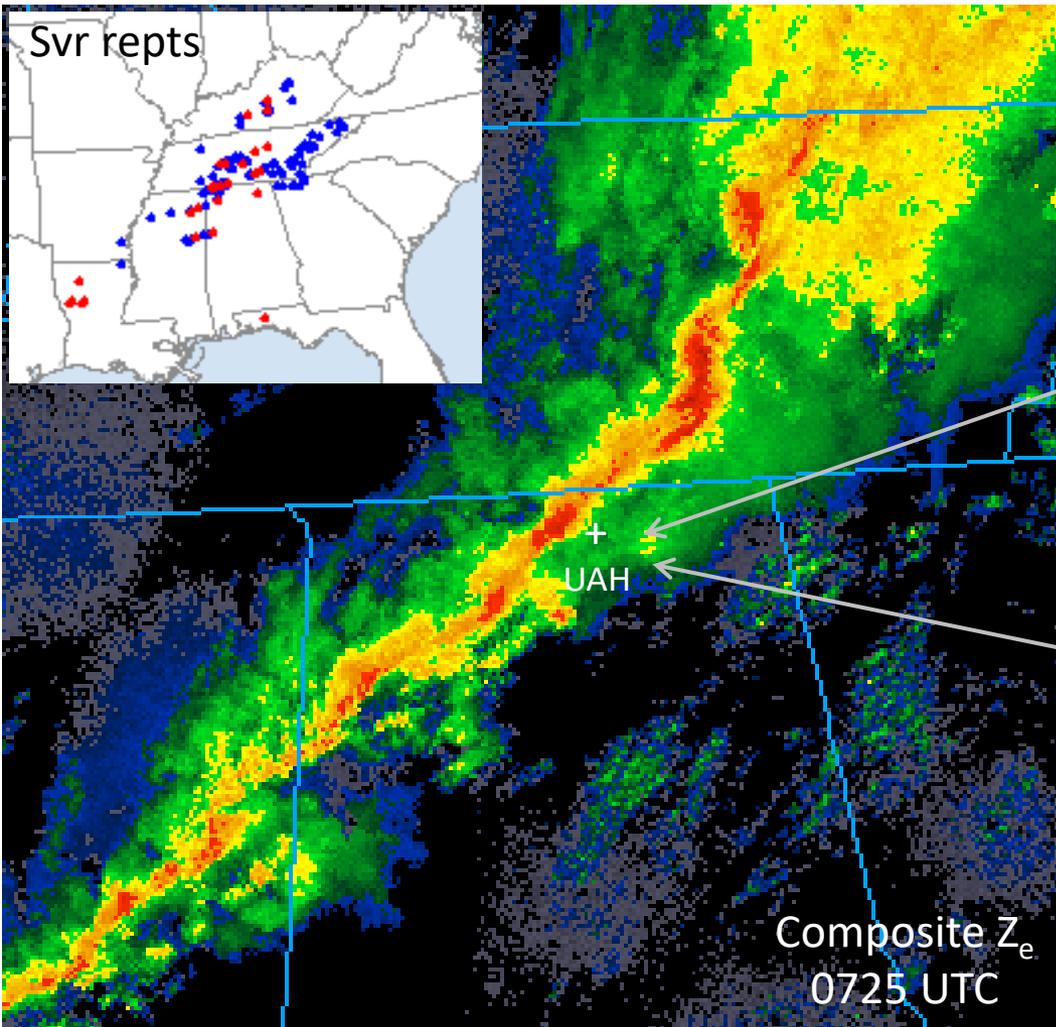
- Improve understanding of the pre-QLCS boundary layer (BL) and its variability (BLV) (primarily Nov-April)
  - Heterogeneities (large scale): significant horizontal advection
  - Heterogeneities: boundaries, evaporational cooling from stratiform rain & showers, heterogeneous surface, variations in vertical shear ( $z_0$ ), shear instabilities (K-H), BL eddies/streaks
- Key question: How do variations in wind/thermo profiles in the pre-QLCS environment produce changes in QLCS internal flows?



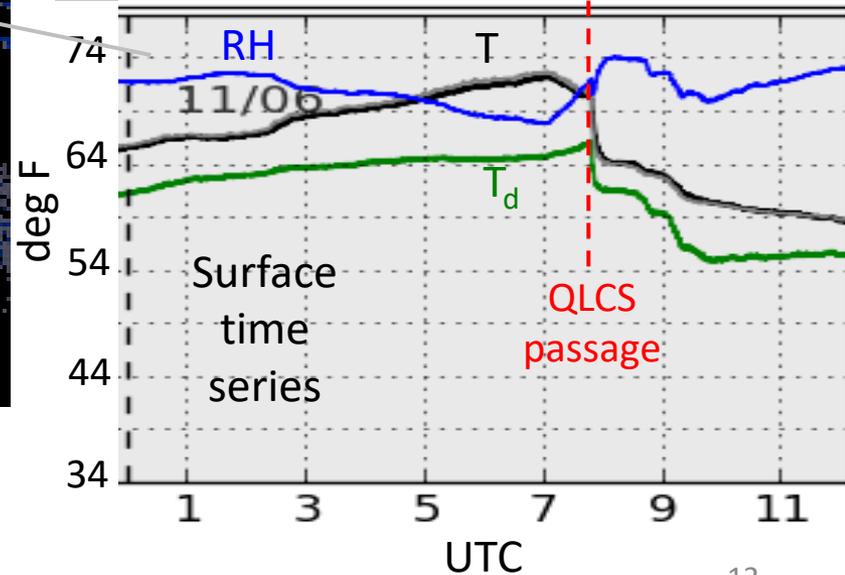
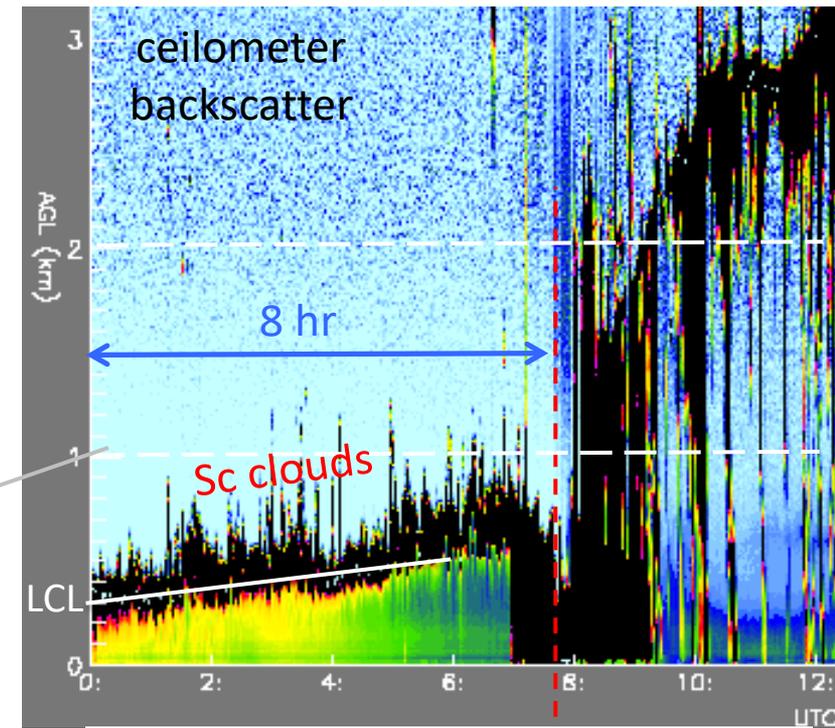


## 2.1 Sc clouds preceding the QLCS

Example: Severe QLCS, 6 Nov 2018



Co-located ceilometer and surface



Stratiform and convective rain preceding the QLCS may play an important role in stabilizing the BL and enhancing wind shear on  $\sim 10$  km scales

# Characteristics of Sc clouds preceding severe cold season QLCS's

## Motivation

- Sc clouds are:
  - An important component of the BL "system"
  - Modulate subcloud static stability/shear
- Sc cloud base height and fraction may serve as a good metric for numerical simulation fidelity

## Hypothesis

- Not a traditional cloud-topped mixed layer
- BL is statically stable in most cases

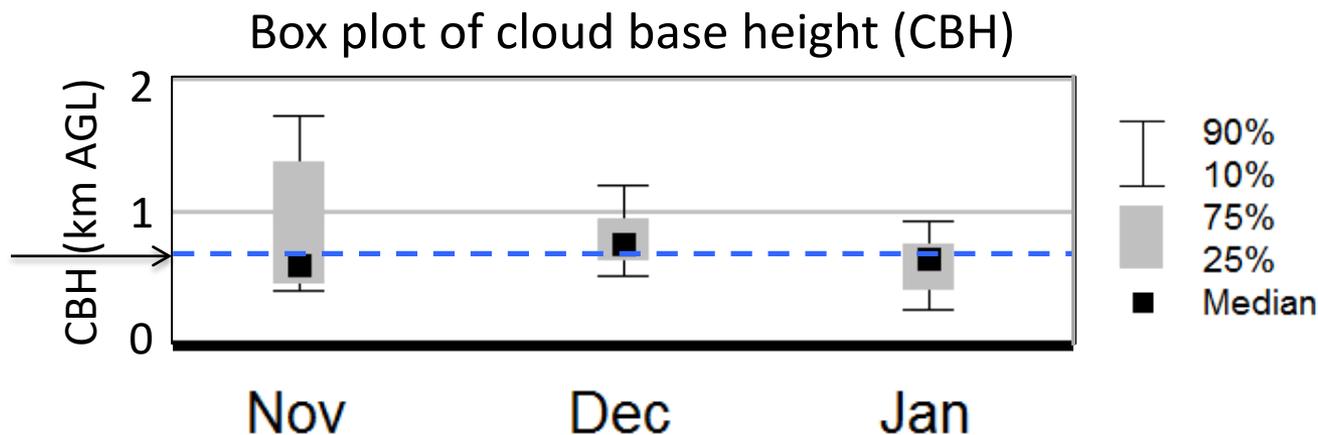
CAMS do not reproduce this cloud-topped PBL to sufficient detail.

Low cloud fraction vs. month  
Based on ceilometer data

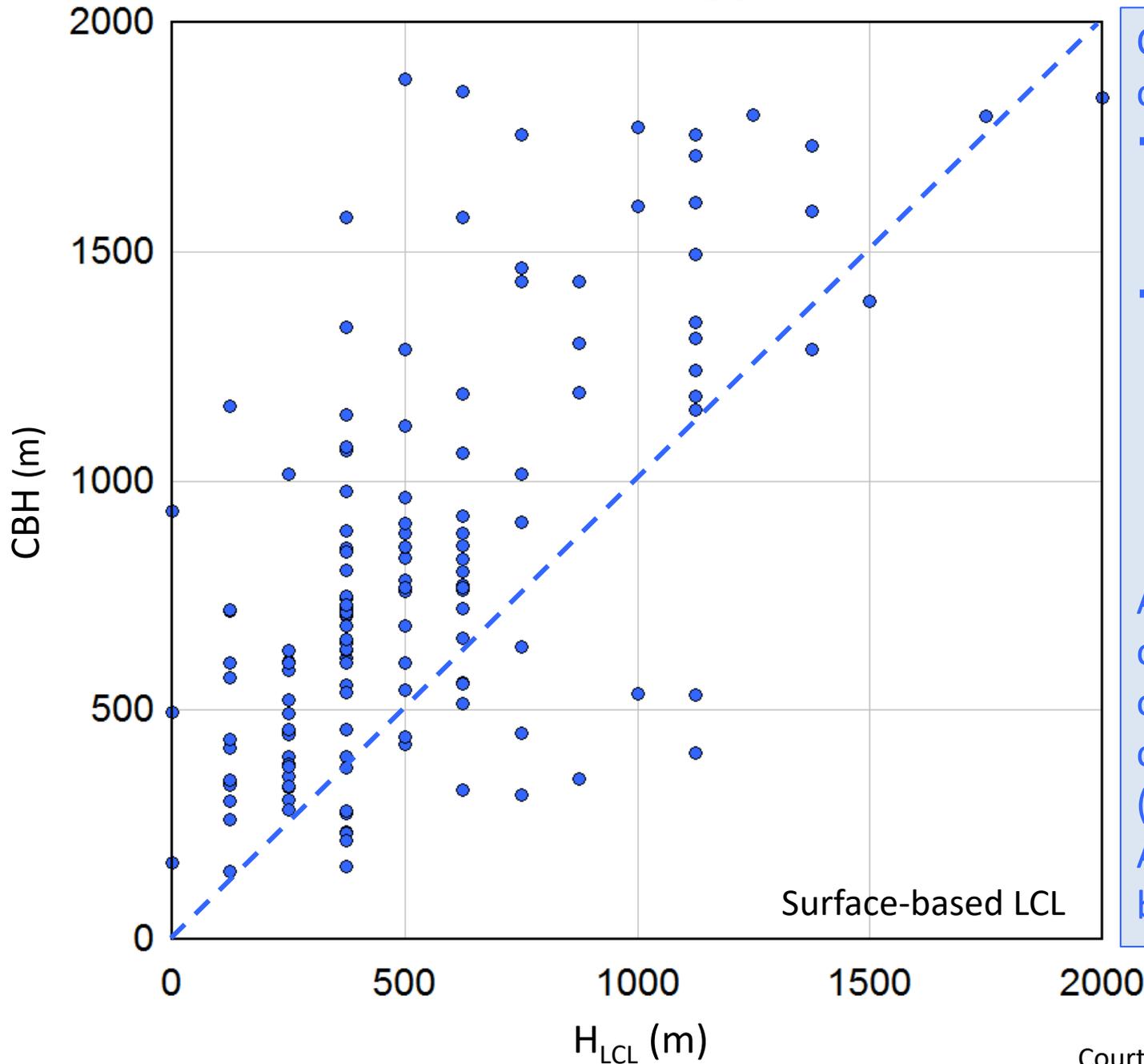
Month	cases	CBH < 2 km (%)
Nov	9	82
Dec	8	89
Jan	7	87
Feb	12	78
<b>Total</b>	<b>36</b>	<b>83</b>

- *Cloud fraction is high*

- *Median CBH is around 700 m AGL*



CBH (ceilometer) vs.  $H_{LCL}$  (sfc data)



CBH >  $H_{LCL}$  in most cases  
→ BL is not mixed: It is statically stable in most cases  
→ Static stability moderates TKE (mixing), and promotes greater wind shear

AMF3 would provide comprehensive data on Sc cloud characteristics (cloud radars, lidars, AERI, scanning C-band radar)

# A future project (planning stages) to examine AV/BLV preceding QLCSs

- *Characterization of Atmospheric Variability, Evolution, And Thermodynamics preceding QLCSs (CAVEAT-Q)*
- PI's: Knupp, B. Geerts (UW), M. Biggerstaff (OU) and Z. Wang (CU)
- Key platform: UW King Air with the following instruments:
  - Compact Raman Lidar (T and w.v. mixing ratio profiles)
  - Wyoming Cloud Radar (W-band, cloud depth, kinematics)
  - Ka-band zenith/nadir profiling radar (KPR) (precip distribution, kinematics)
  - In situ
- Other facilities: all UAH, two mobile SR radars from OU, maybe others
- Objectives:
  - Advance understanding of BL characteristics, spatial variability, and driving physical processes in the pre-QLCS environment
  - Understand how variations in pre-QLCS BL structure, in particular low-level shear and CAPE, affect storm-scale process within QLCSs, especially the evolution of vorticity within or near shallow convective updrafts
- Likely time period: Feb – Apr 2022 or 2023 (cold season to spring conditions)
- Domains: near Courtland (Alabama) and Mississippi delta (TBD)
- Additional objectives (e.g., aerosol) would be easy to incorporate

# A future project (planning stages) to examine AV/BLV preceding QLCs

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**Ideal example of synergism with AMF3 capabilities (wind, thermodynamic, and cloud profiling) and ARM/ASR science**

# Summary

- Challenges in network design
  - trees, orography
  - document both BL and cloud-scale flows
- Complex surfaces (turbulent fluxes) introduce heterogeneities on small scales that likely influence deep convection
- Mesoscale variability is important to document
  - Time-to-space conversion of time series data (profilers)
  - Scanning radar
  - Aircraft measurements if/when available

# Backup slides

# ***BL heterogeneity***

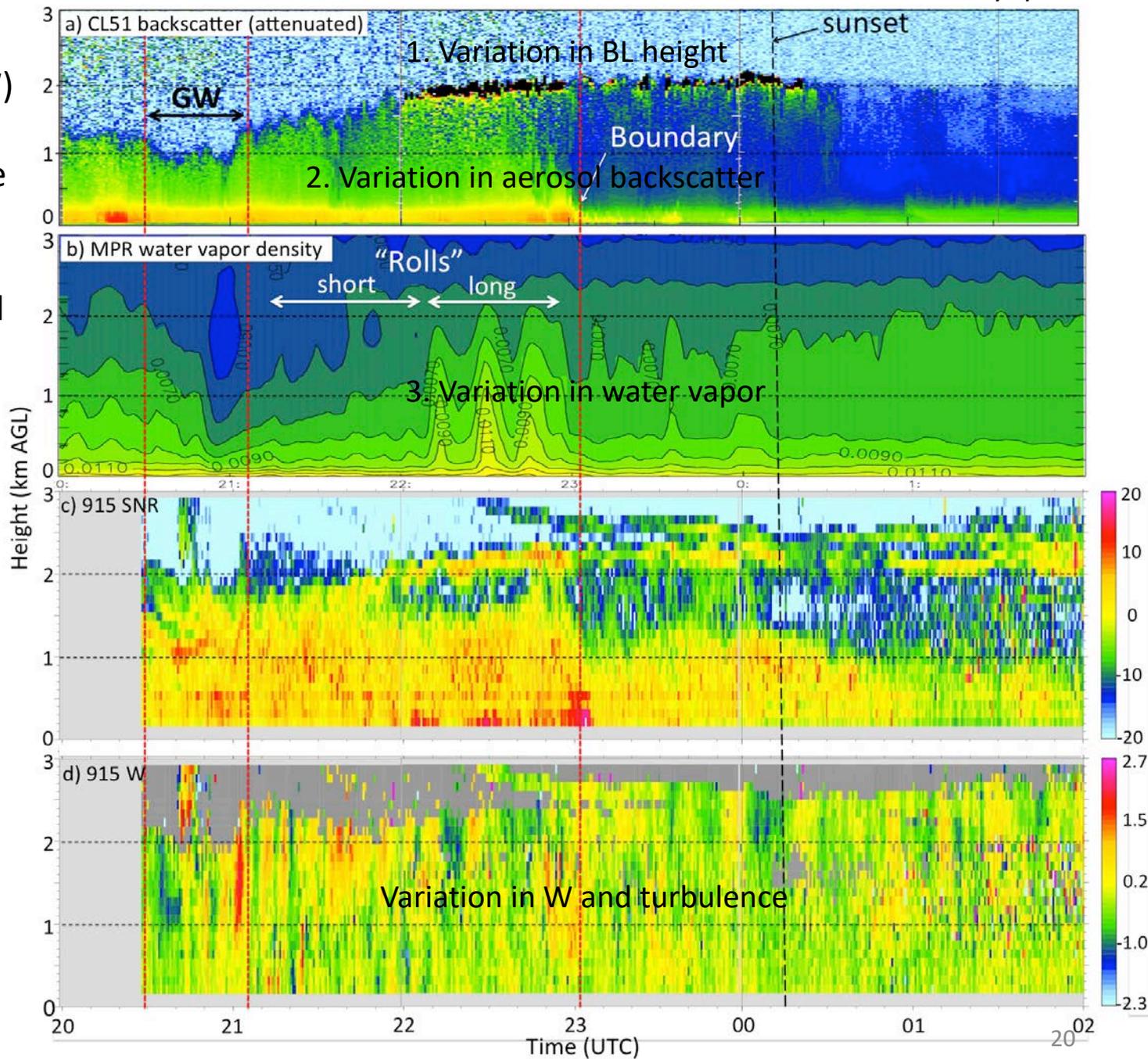
- Impact on tornadic QLCs and supercell storms
- Importance in convective initiation (CI)
- Modulation of Cu cloud fields
- Mesoscale variations in surface properties ( $z_0$ , fluxes)
- Example:
  - 4/9/13
  - Wheeler lake breeze

Note: Relative forcing by mesoscale heterogeneities is likely a more important regulator on deep convection during the warm season when synoptic-scale forcing is small and surface fluxes are large.

## MIPS observations

- a) Gravity wave (GW)
- b) BL rolls
- c) Boundary passage

Appreciable variation  
in  $z_i$ , water vapor, and  
relative aerosol  
loading.

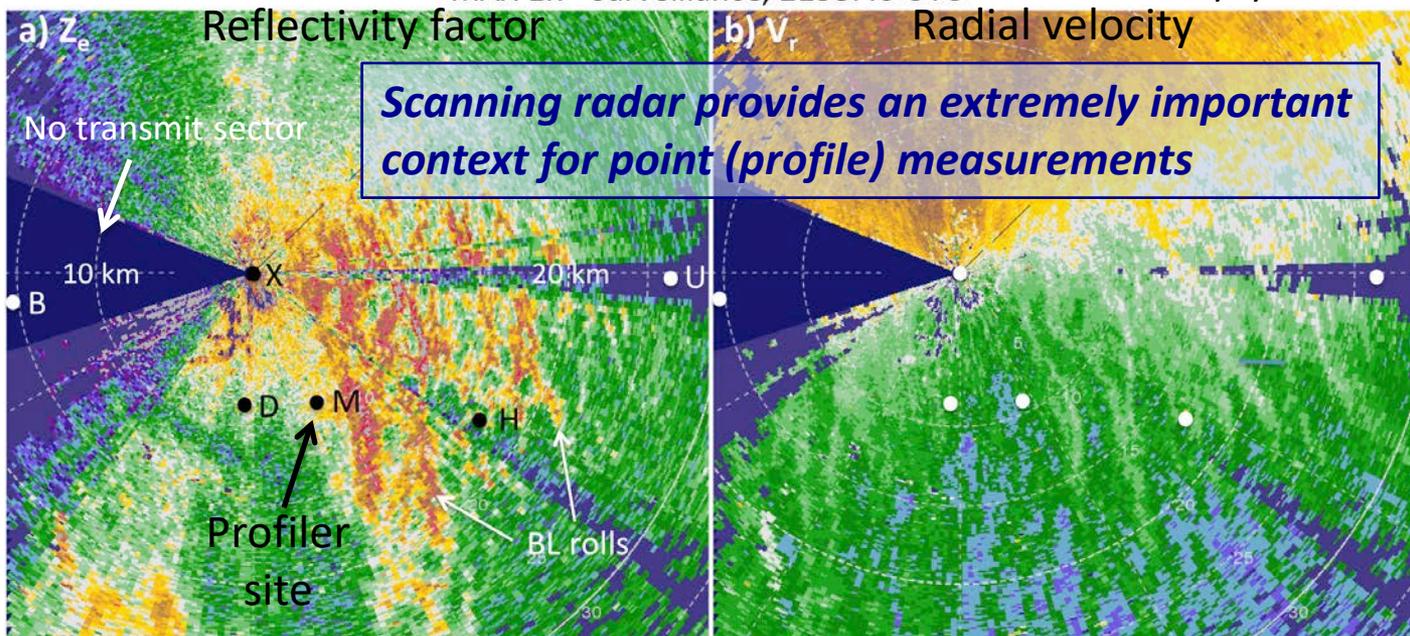


# MAX Observations

MAX 1.7° surveillance, 2138:40 UTC

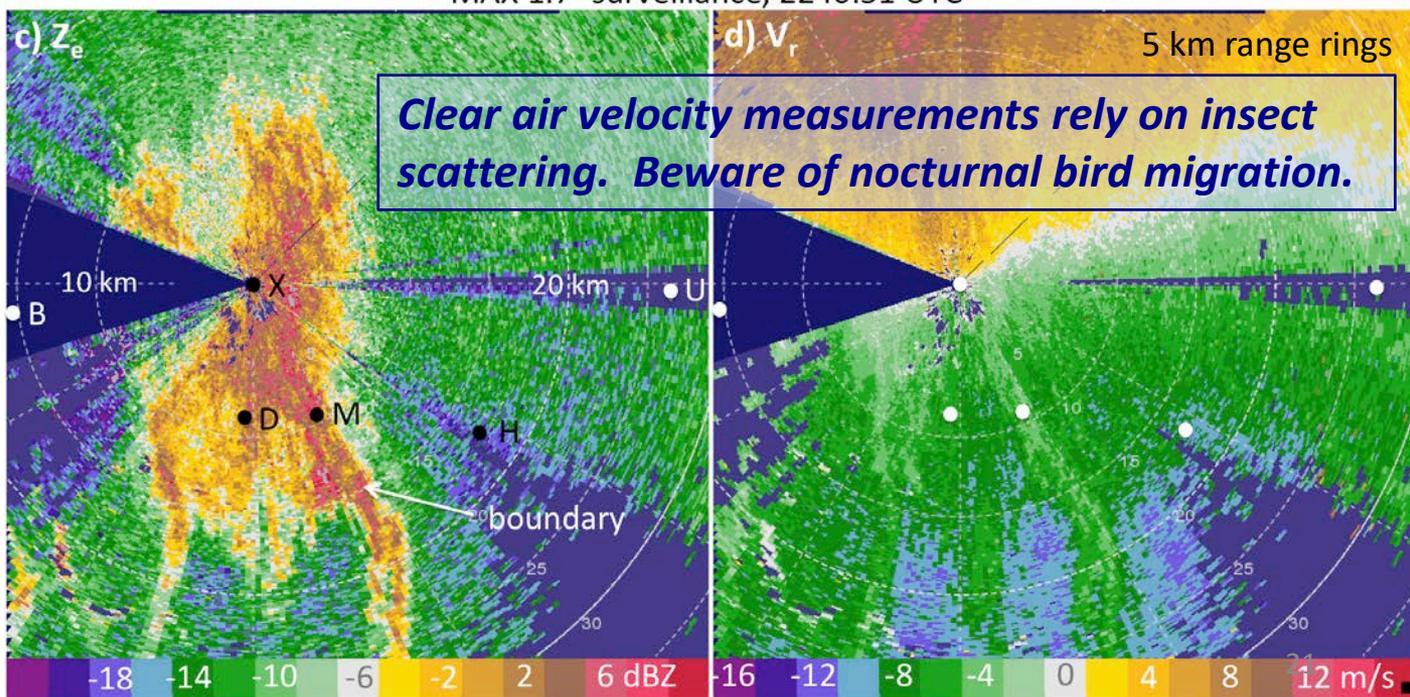
4/9/2013

2139 UTC (top):  
Limited region of enhanced  $Z_e$  and  $V_r$  streaks, consistent with BL rolls.  
Distinct difference in  $T$  &  $T_d$  between KDCU (D) and KHSV (H)



MAX 1.7° surveillance, 2246:51 UTC

2247 UTC (bottom):  
Presence of boundary (not a lake breeze), superimposed within a region of bands of relatively enhanced  $V_r$



- M: MIPS
- X: MAX
- D: KDCU ASOS
- H: KHSV ASOS
- U: UAH