

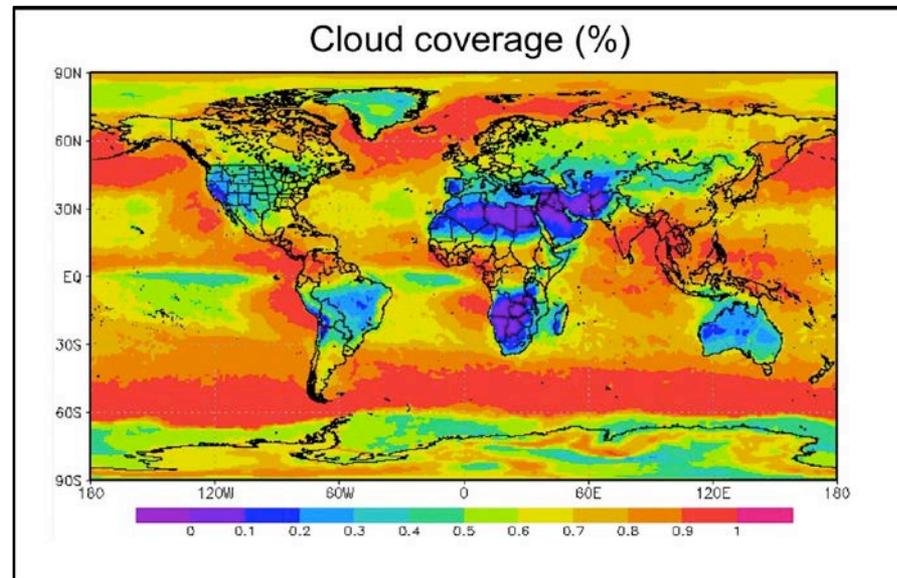
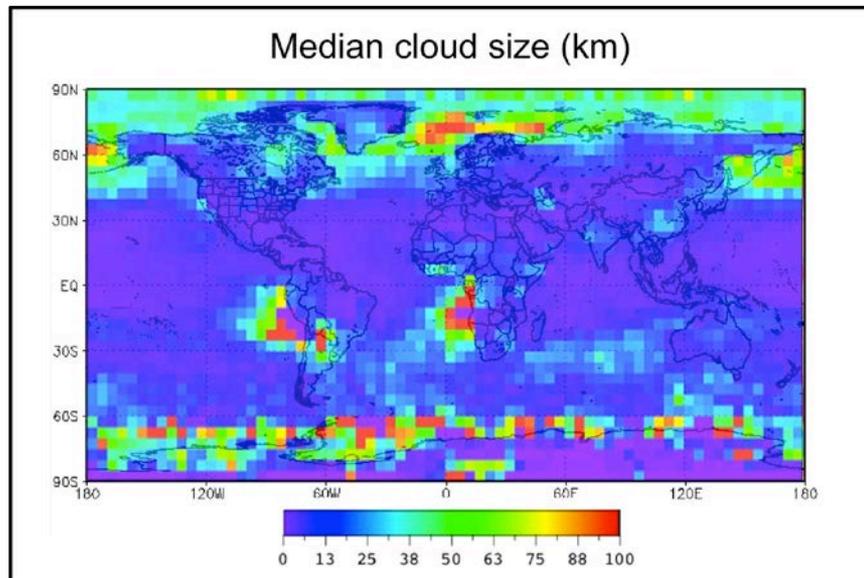
# Multiscale Methods for Accurate, Efficient, and Scale-Aware Models of the Earth System

**Principal Investigator: Bill Collins**

## **Science Team Leads and Institute Liaisons:**

- **Atmosphere** Steve Ghan
- **Ocean** Todd Ringler
- **Computational Science** Carol Woodward
- **Multiscale UQ** Don Lucas
- **FASTMath Liaison** Carol Woodward
- **SUPER Liaisons** Lenny Oliker and Sam Williams
- **QUEST Liaison** Bert Debusschere

# Clouds scales beyond current model resolution



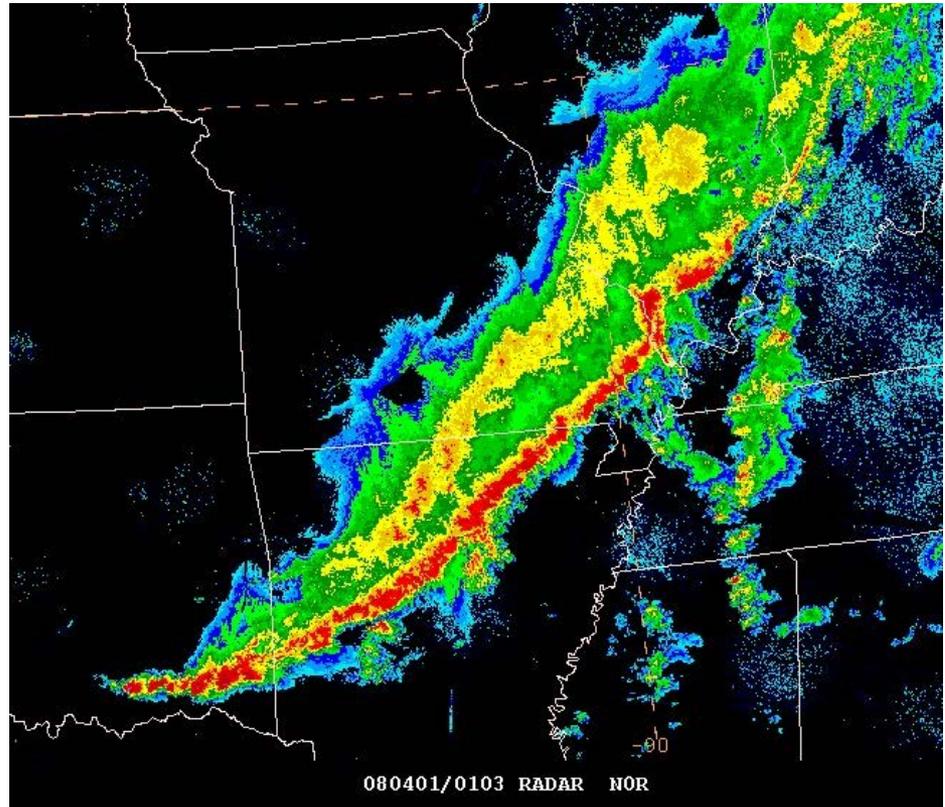
↑↑    ↑    ↑  
CSSEF UHR    AR5    AR4

- Many climate-relevant features are localized in space and time
- Features can be: stationary, move at advected speeds, or evolve from wave dynamics
- Refinement can greatly reduce computational resources needed

Examples:

- Tropical cyclones
- Squall lines / storm fronts
- Atmospheric rivers
- Ocean eddies

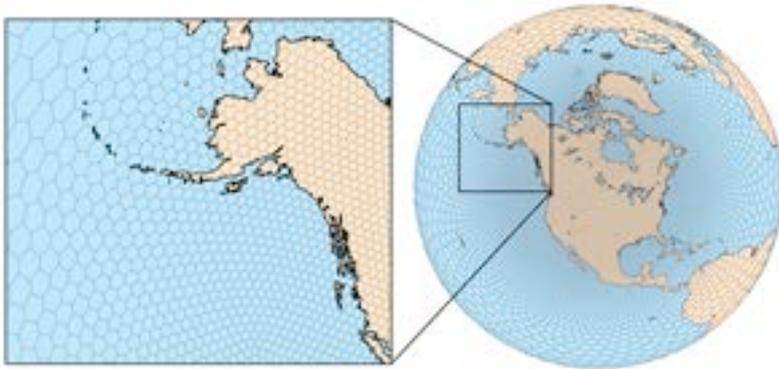
→ Refine where/when required



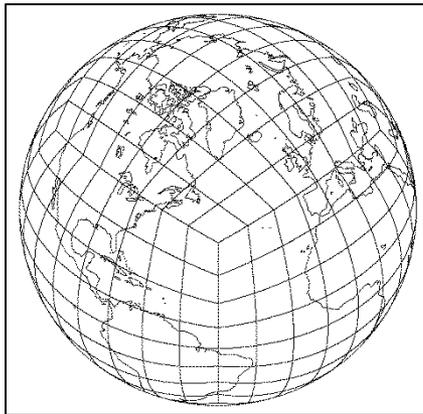
Radar image of a squall line over central US

## Variable Mesh Dycores

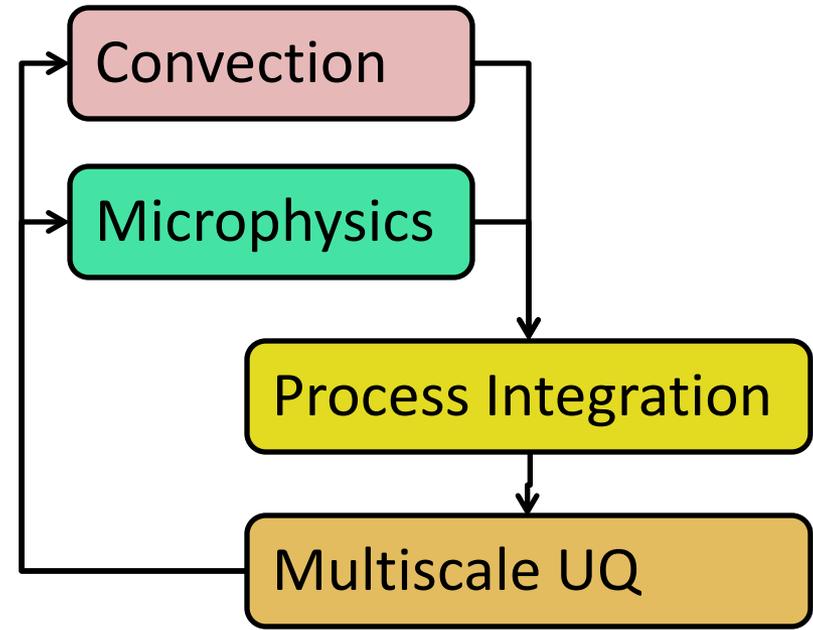
Model for Prediction Across Scales (MPAS)



Spectral Element Dycore



Physics-Dynamics Interface

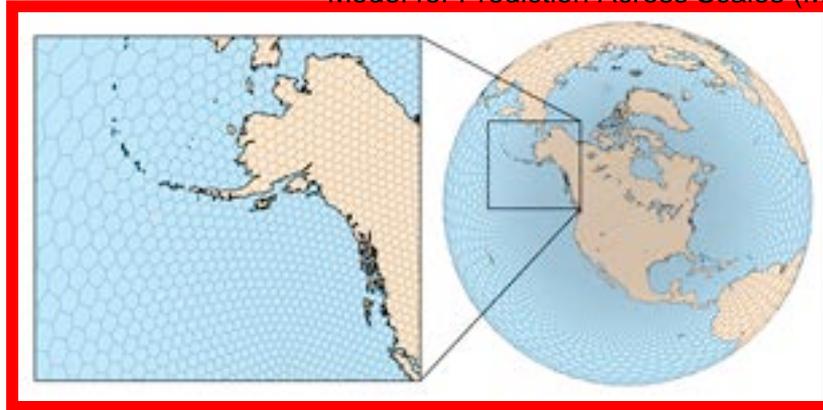


Atmosphere  
Ocean

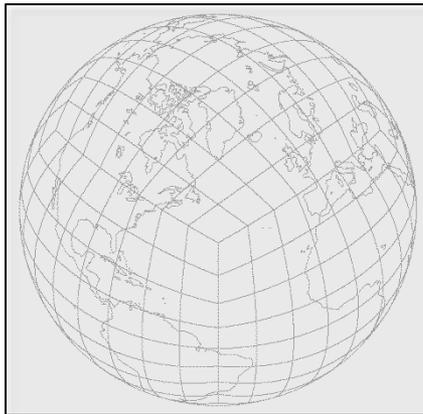
Mesoscale Eddy  
Treatments

## Variable Mesh Dycores

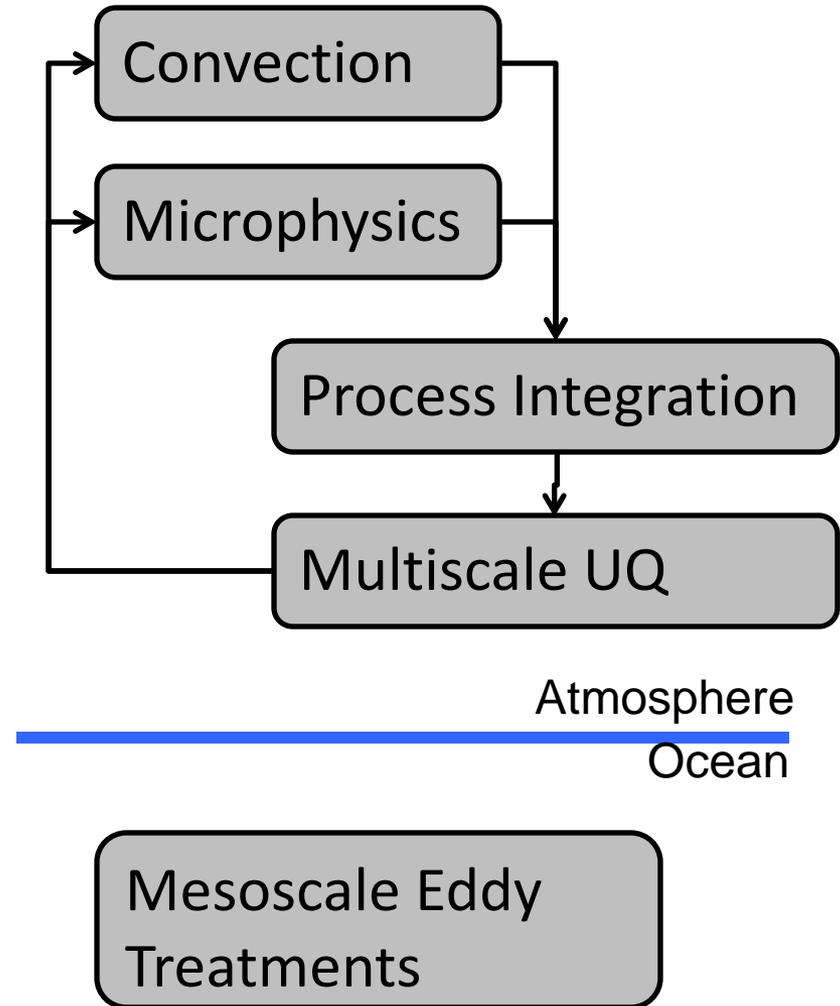
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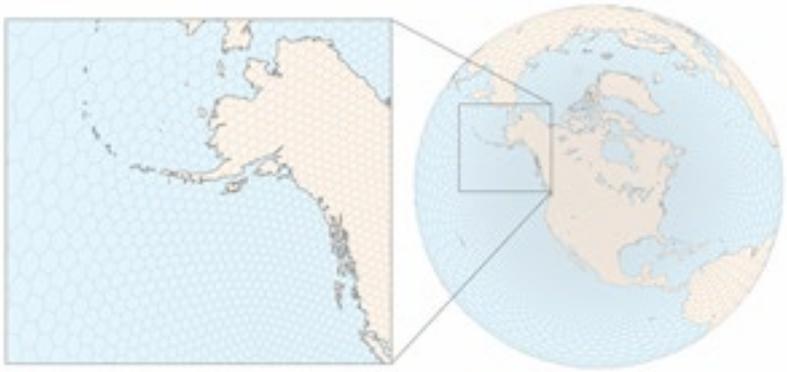


Physics-Dynamics Interface



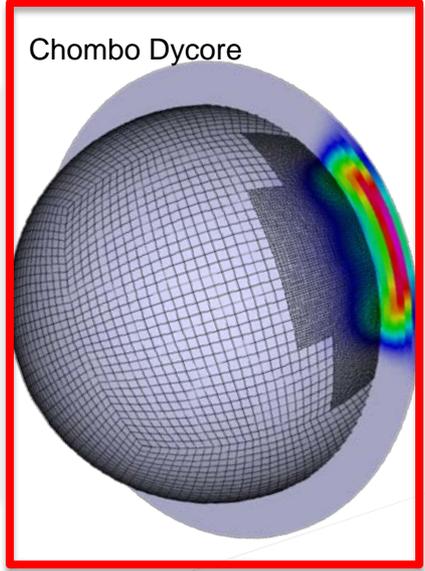
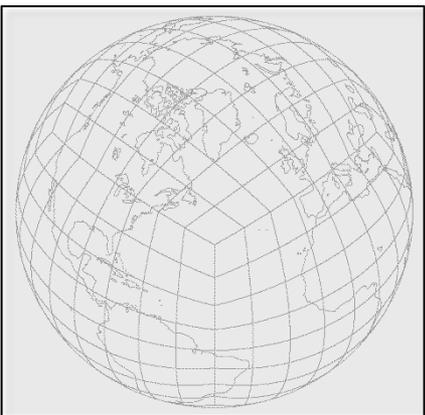
## Variable Mesh Dycores

Model for Prediction Across Scales (MPAS)

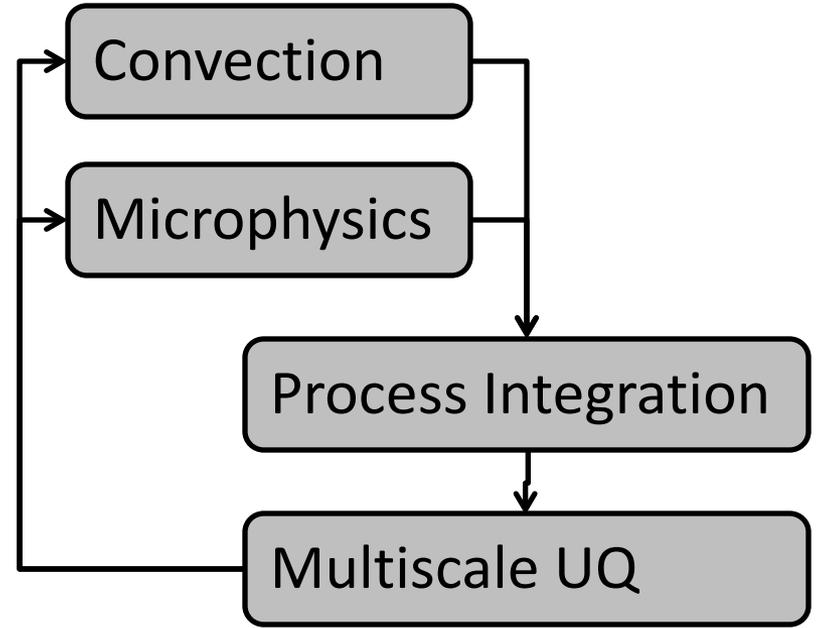


Spectral Element Dycore

Chombo Dycore



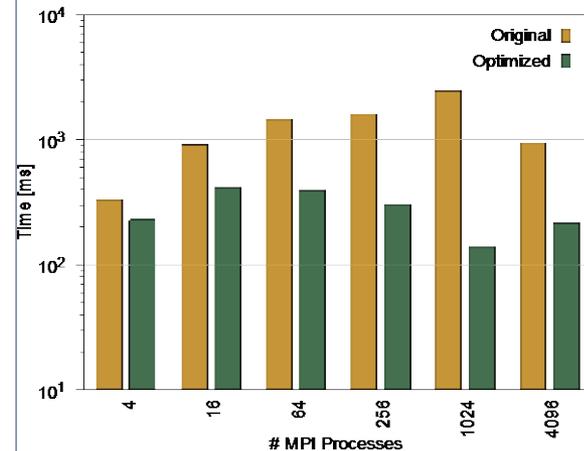
Physics-Dynamics Interface



Atmosphere  
Ocean

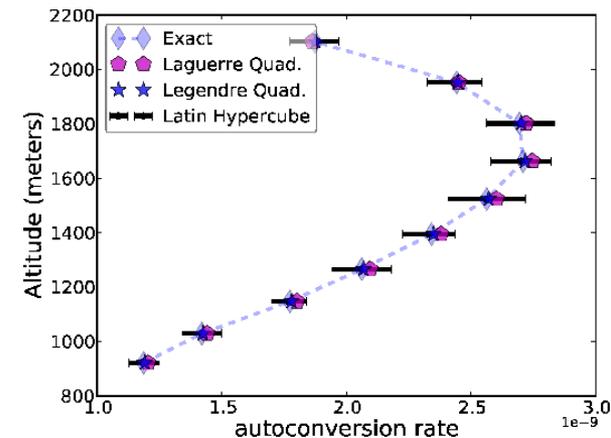
Mesoscale Eddy  
Treatments

- **SUPER:** Investigating intra-node optimization of MPAS-Ocean via Space-Filling-Curve ordering and inter-node optimization via aggregate communication schemes
- **FASTMath:** Developing preconditioners for scalable implicit solves for CAM-SE
  - Journal paper submitted
  - Initiated multilevel solver development
- **FASTMath:** Developing Implicit-Explicit Non-hydrostatic AMR Dycore; Completed tests for SWE, Linearized Gas Dynamics
- **QUEST:** Developed and applied a deterministic quadrature approach for integrating microphysics over subgrid variability
  - To CLUBB autoconversion & accretion
  - Journal paper submitted
- **QUEST:** Applying Bayesian Compressed Sensing for surrogate development and sensitivity analysis in CLUBB (29 input params)

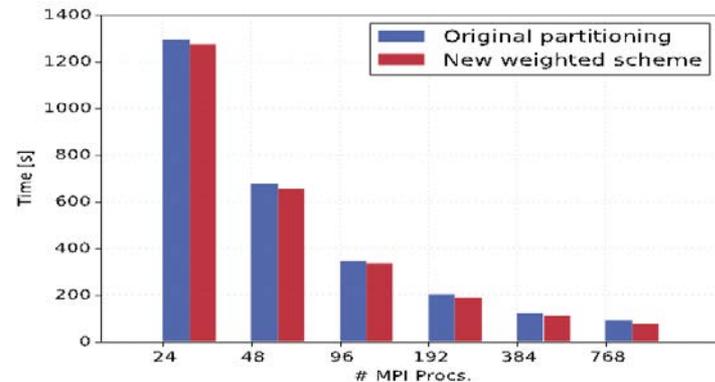


*Optimized communication potential via a micro-benchmark developed to capture the MPAS-O halo exchange (Edison)*

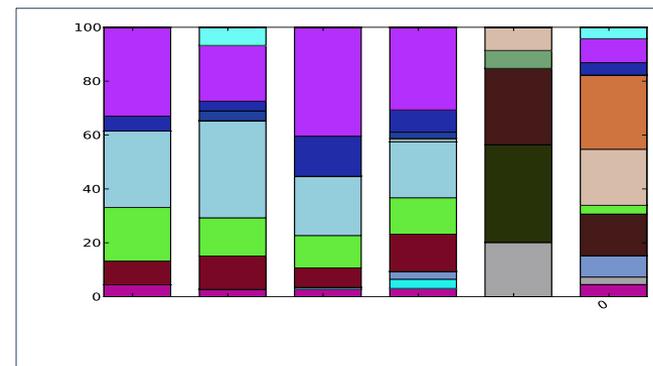
Autoconversion at 72 hours in the RICO case, showing superior accuracy of quadrature results based on just 16 samples over Latin Hypercube results based on 500 random samples



- **SUPER:** New MPAS-O partitioning scheme based on hyper-graph with weights taking into account the computations and communication due to local cells as well as halo cells
- **FASTMath:** Initiated project to analyze splitting methods and convergence with time step for microphysics
- **QUEST:** Bayesian Compressed Sensing to perform sensitivity analysis and build liquid cloud fraction models in CLUBB
  - Adaptive approach generates surrogates with fewer terms and reduced over-fitting



Reduction in total MPAS-O run time using hyper-graph partitioning, with 5% speedup at highest concurrency



BCS derived sensitivity: similar to full 1<sup>st</sup> order basis approximation, with only half as many terms

## Implicit Time Integrators (KE, MT, C. Woodward)

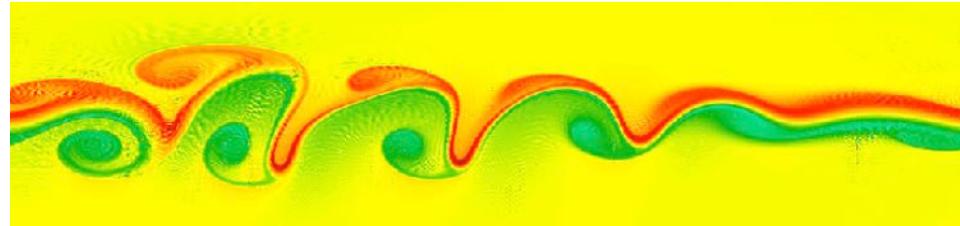
- In the CAM-SE dycore, the hyperviscosity term is now included within the implicit solve of the primitive 3D equations
- Allows a 2nd order accurate implicit solution at the time-scale of tracer advection.
- Optimization is under way; preconditioning options being explored for efficiency.

## Adaptive Mesh Refinement Dycore (HJ, E. English, w/ FASTMath team @ LBNL)

- 2D shallow water complete (1 paper)
- 3D non-hydrostatic dynamics implemented, with 1000:1 aspect ratio (non-linear implicit vertical)
- Running DCMIP tests with space-time AMR

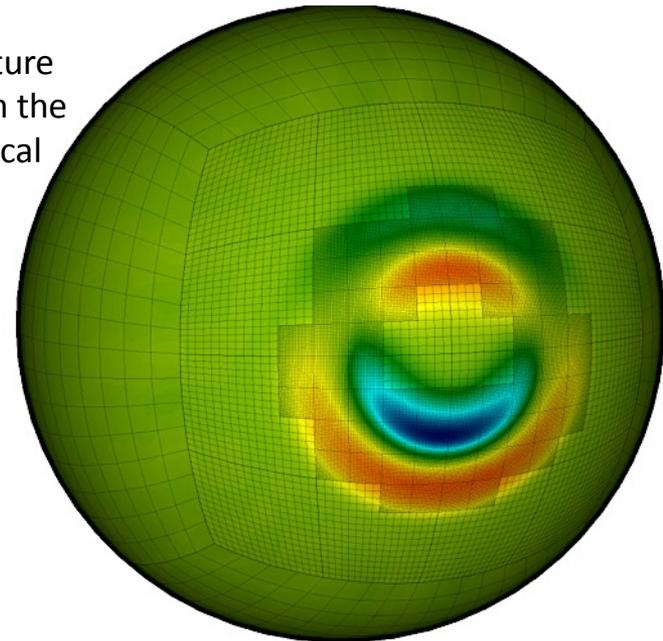
## CAM Multi-resolution techniques (MT)

- Paper (in review, GMD), “The spectral element method on variable resolution grids: Evaluating grid sensitivity and resolution-aware numerical viscosity”
- Presentation (MT) at ICOSAHOM, Salt Lake City, June 2014.



Relative vorticity after 6 days of the Galewsky et al. (2004) baroclinic instability test case, using the implicit method.

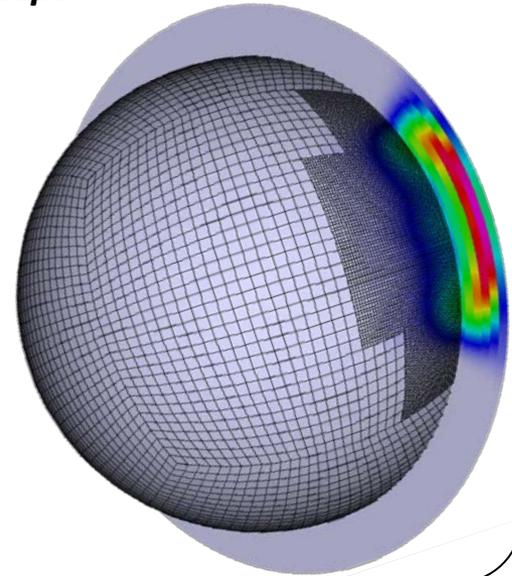
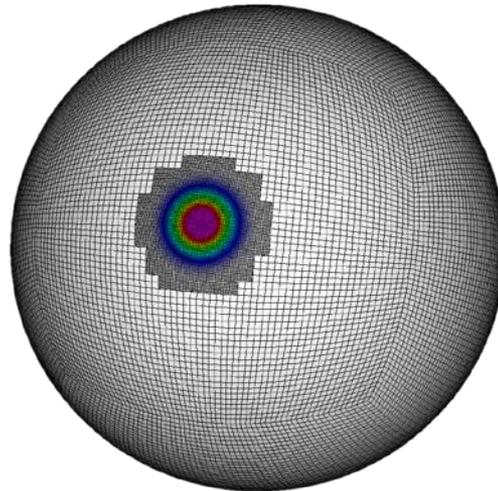
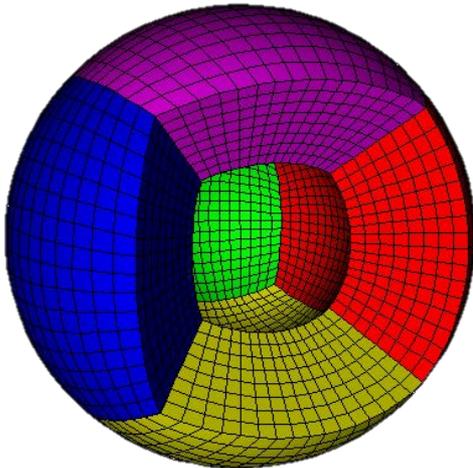
Potential temperature perturbations from the interaction of vertical structure with Coriolis forces. Test uses  $1^\circ$  mesh with 16x AMR in the new 3D non-hydrostatic dycore.



# Goals of dynamic AMR atmospheric simulations

Create a high-resolution, high accuracy dycore to eliminate local numerical errors

- Fully non-hydrostatic dynamics (4<sup>th</sup>-order in space and time)
- Excellent scaling (100k+ cores) without vertical explicit acoustic CFL limitation
- Anisotropic adaptive mesh refinement in space, refinement in time
  - 3 levels with 4x refinement produces > **5 more digits** of accuracy
  - Starting with 1° base level, **resolution < 500m, ~2s time steps**



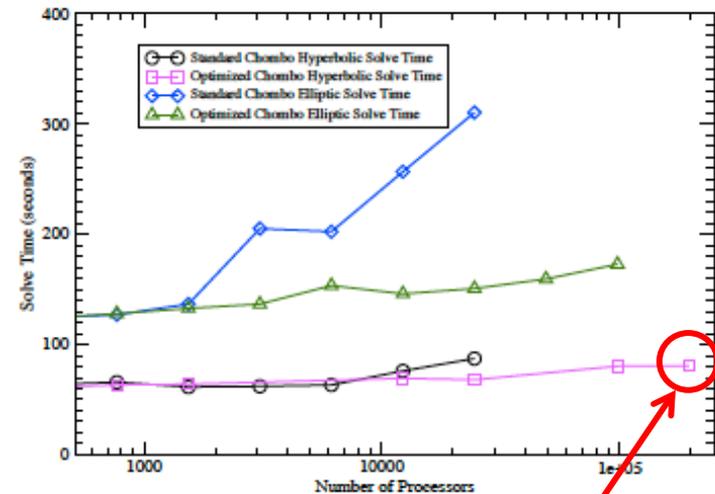
# Chombo AMR code is well-established, scalable

## Broad set of applications leveraging Chombo framework with complex physics:

- MHD for tokamaks (R. Samtaney, PPPL)
- Cosmology: CFD + particles (Miniati, ETH)
- Space plasmas: compressible CFD electromagnetic, kinetic effects (G. Zank, UA)
- Astro. MHD turbulence (McKee / Klein, UCB)
- SF Bay Hydrology modelling (CA DWR)
- Microscale fluids (UNC / UCD / LBNL EFRC)
- Heat transfer in nuclear reactors (LBNL)
- Nuclear reactor safety - hydrogen combustion entrainment models (Calhoun, CEA-Saclay)
- Type II Supernovae (Woosley, UCSC, LLNL)
- 4D gyrokinetic tokamak edge plasmas (LLNL) ...

Long-term ASCR investment: SciDAC, FastMATH ...

→ Target platforms extreme-scale / DoE LCF's



**200k processor benchmark,  
nearly ideal weak scaling  
on similar hyperbolic problems!  
[EuroPar2011]**

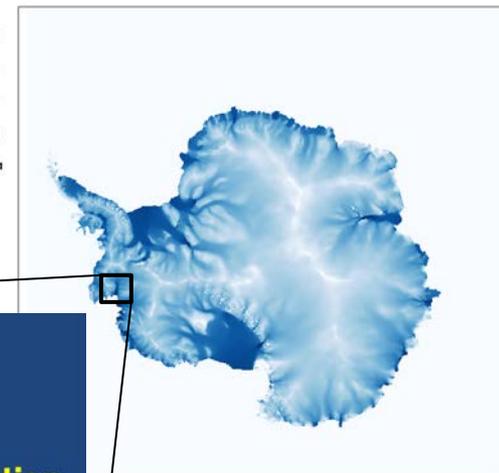
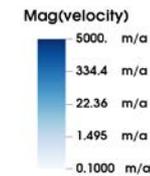
## Objectives

- Understanding ice sheet dynamics is crucial for credible predictions of future sea level rise (SLR)
- Very fine resolution (better than 1 km) needed to resolve dynamics of grounding lines – unfeasible for entire ice sheets
- Dynamic Adaptive Mesh Refinement (AMR) brings fine resolution to bear only where needed to resolve the dynamics.

## Impact

- Enables modeling grounding line dynamics of marine ice sheets with sufficient spatial resolution to correctly model advance or retreat.
- AMR models dynamics correctly (vs. lo-res reduced or no mobility)
- Enables continental-scale modeling of ice sheets at the resolution / computational cost for global climate models (GCMs).

(below) Computed ice velocity for Antarctica, mesh and grounding line for Pine Island Glacier.



## Amundsen Embayment Ice Sheet Simulation

One possible climate scenario (Payne et al.)  
simulated using SciDAC-funded BISICLES code



Office of  
Science

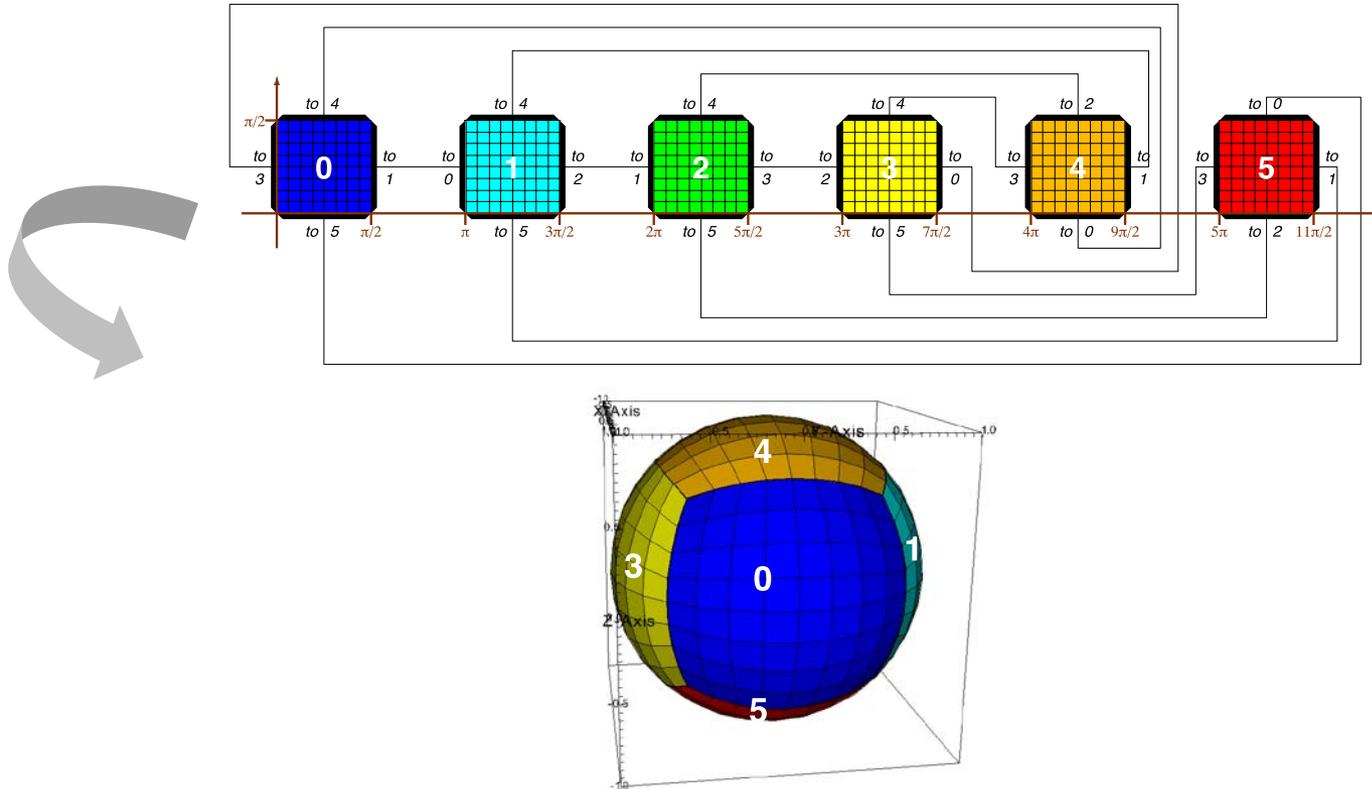


Los Alamos  
NATIONAL LABORATORY



# Cubed Sphere Formulation

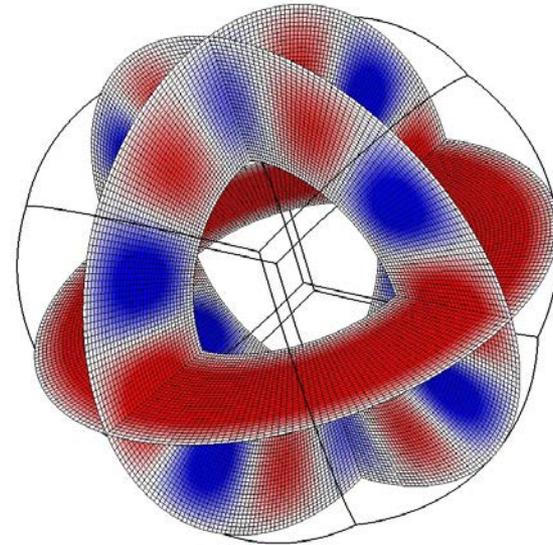
- Use a “multi-block” mapping between rectangular index space and CS panels
- Use an arbitrary stretched grid in the radial direction (pressure coordinate, etc.)



# Conservative FV on Cubed Sphere “Shells”

- Riemannian form of compressible Euler [Ullrich Jabl. JCP 2012]
- Conservative, in flux divergence form, with “~” indicating deviation from constant hydrostatic background state
- Cubed sphere metrics calculated analytically, mapped to unit sphere
- “Shells” because we neglect the radial metric dependencies

$$\begin{array}{ll} \text{density} & \frac{\partial J \tilde{\rho}}{\partial t} = \dots \\ \text{potential temp.} & \frac{\partial J \tilde{\Theta}}{\partial t} = \dots \\ \text{humidity} & \frac{\partial J \tilde{Q}}{\partial t} = \dots \\ \text{momentum} & \frac{\partial J \rho \mathbf{u}}{\partial t} = \dots \end{array}$$



$$\mathbf{u} = u^\alpha \mathbf{g}_\alpha + u^\beta \mathbf{g}_\beta + u^\xi \mathbf{g}_\xi$$

- At 1° resolution, lower atmosphere cells can have ~1000:1 aspect ratio
- Euler has (fast) acoustic waves, (almost as fast) gravity waves

- Sound wave-based CFL for explicit methods → 1000x smaller time step
- Implicit methods used in place of explicit time integration in the vertical direction.

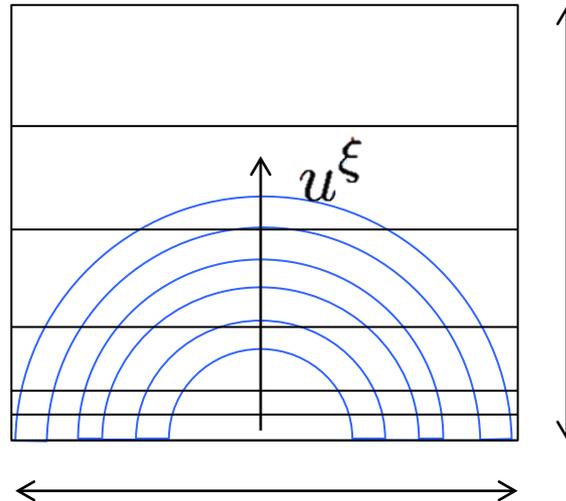
$$\frac{\partial J\tilde{\rho}}{\partial t} + \frac{\partial}{\partial \xi}(J\tilde{\rho}u^\xi) = \dots$$

$$\frac{\partial J\tilde{\Theta}}{\partial t} + \frac{\partial}{\partial \xi}(J\tilde{\Theta}u^\xi) = \dots$$

$$\frac{\partial J\tilde{\rho}u^\xi}{\partial t} + \frac{\partial}{\partial \xi}(JG^{\xi\xi}\tilde{p}) = \dots$$

$$\frac{\partial J\tilde{p}}{\partial t} + \gamma J\tilde{p}\frac{\partial u^\xi}{\partial \xi} = \dots$$

$$p = p_0 \left( \frac{R_d \rho \theta}{p_0} \right)^\gamma$$

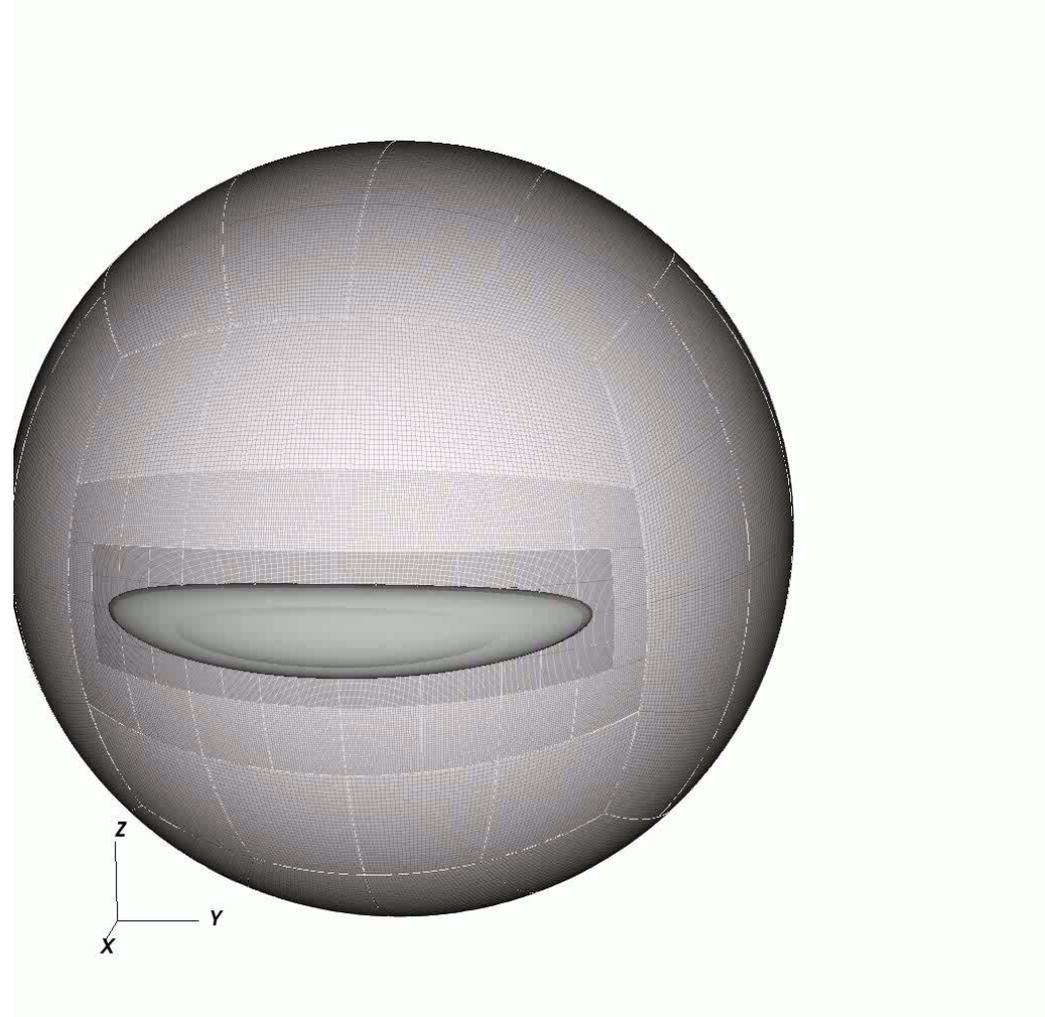


Vertical acoustic waves cross thin cells too quickly for explicit CFL!

Horizontal acoustic waves ok within a wide aspect ratio cell

# Demonstration Calculation with Refinement

- **Dynamics:** Prescribed Hadley circulation
- **Physics:** Large-scale condensation
- **Lower boundary:** Aquaplanet
- **H2O processes:** Advection & condensation
- **Movie:** 3D vapor and surface rainfall
- **Simulation time:** 30 days
- **AMR grid:** 0.7° resolution,  
3 mesh tiers

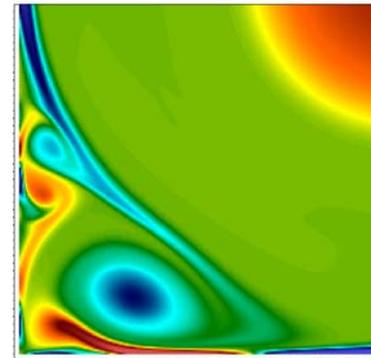


What are we hoping to accomplish?

- 4<sup>th</sup>-order in space and time, without limitation of vertical explicit CFL
- Near-perfect scaling (CHOMBO explicit implementations have gone to 100k+)
- Anisotropic adaptive mesh refinement, 4<sup>th</sup>-order in space, refinement in time
  - 4 levels of 4x refinement produces almost **5 more digits** of accuracy
  - 1° base level, that would be < **500m resolution**, not feasible with uniform grid

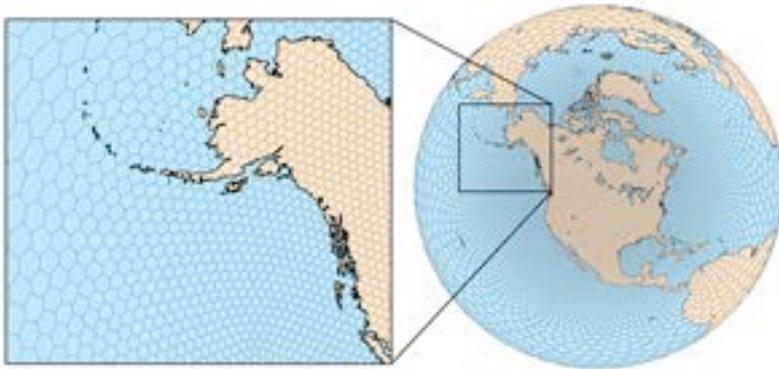
Future plans:

- Finish non-hydrostatic implementation,
- Add orography
- Integrate with various “column physics” to observe refinement behavior
- Support “grid insensitive,” time-accurate physics parameterizations

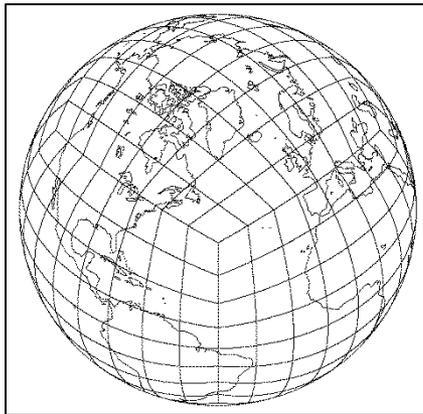


## Variable Mesh Dycores

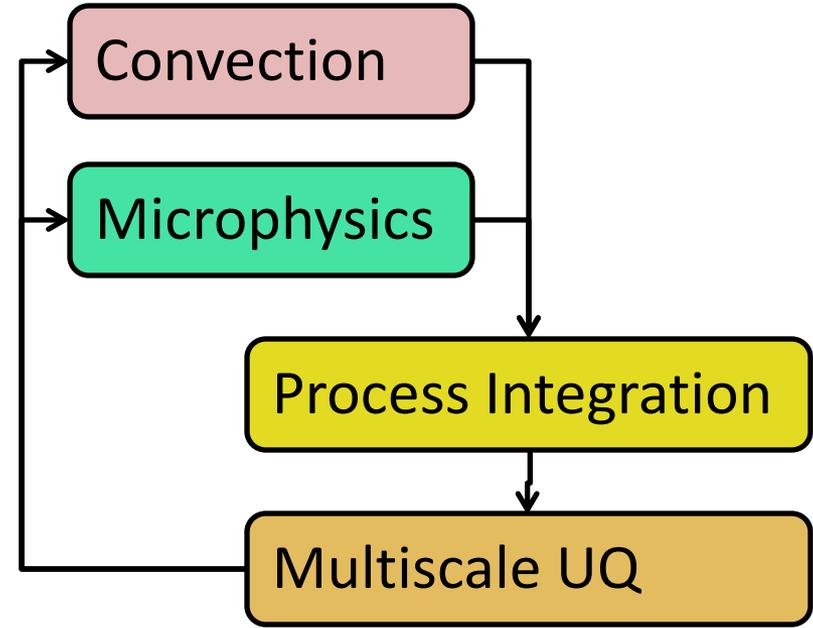
Model for Prediction Across Scales (MPAS)



Spectral Element Dycore



Physics-Dynamics Interface

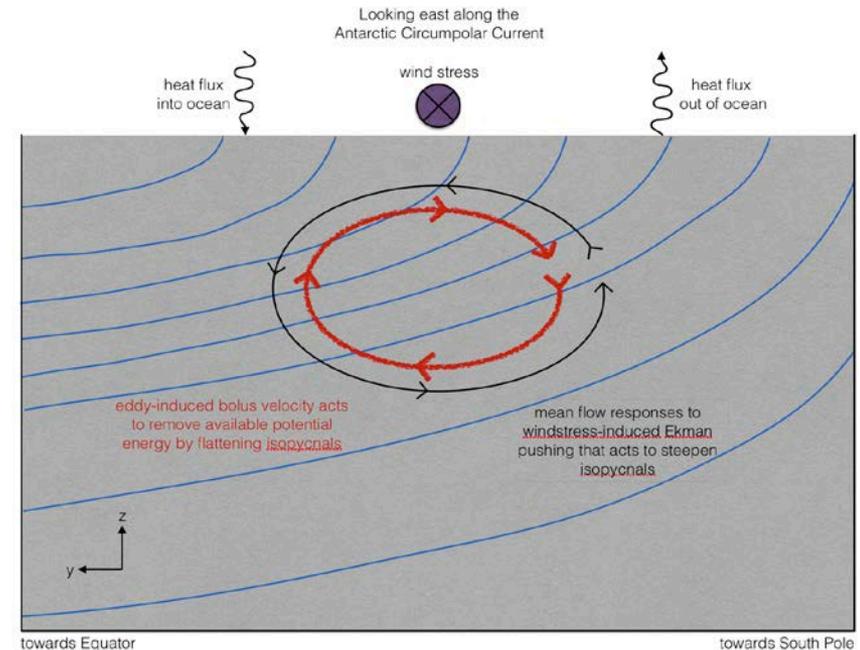


Atmosphere  
Ocean

Mesoscale Eddy  
Treatments

# New multiscale-capable ocean eddy schemes

- Since the seminal work of Gent and McWilliams (GM, 1990), ocean models have used two (typically opposing) velocities: mean and bolus.
- This has caused confusion, ambiguity and inconsistencies in how we characterize the fluid velocity, as well as hinder parameterization development.
- Based on recent theoretical analysis by Young (2012), we have re-expressed the MPAS-O governing equations to solve for the total (residual) velocity.
- A direct comparison to GM is underway.

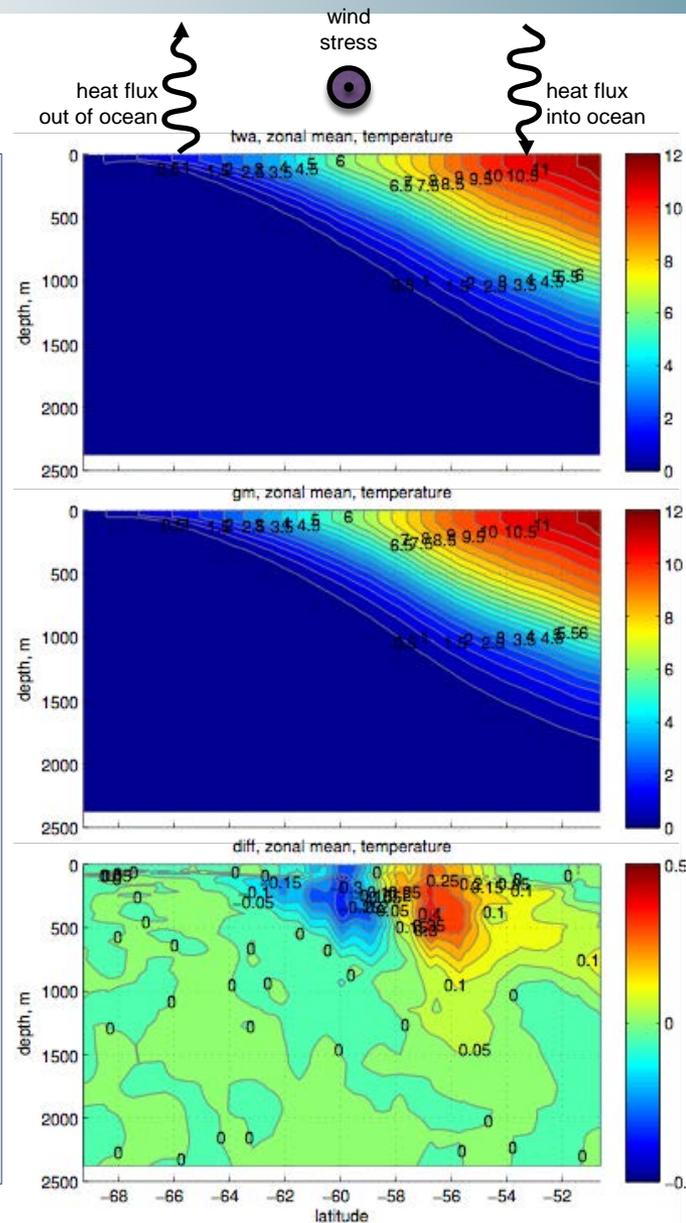


The dynamically-relevant, residual velocity can be solved for directly by recasting the mesoscale eddy parameterization as a force in the momentum equation.

Mesoscale Eddy  
Treatments

# Residual-Mean Scheme for Mesoscale Eddies

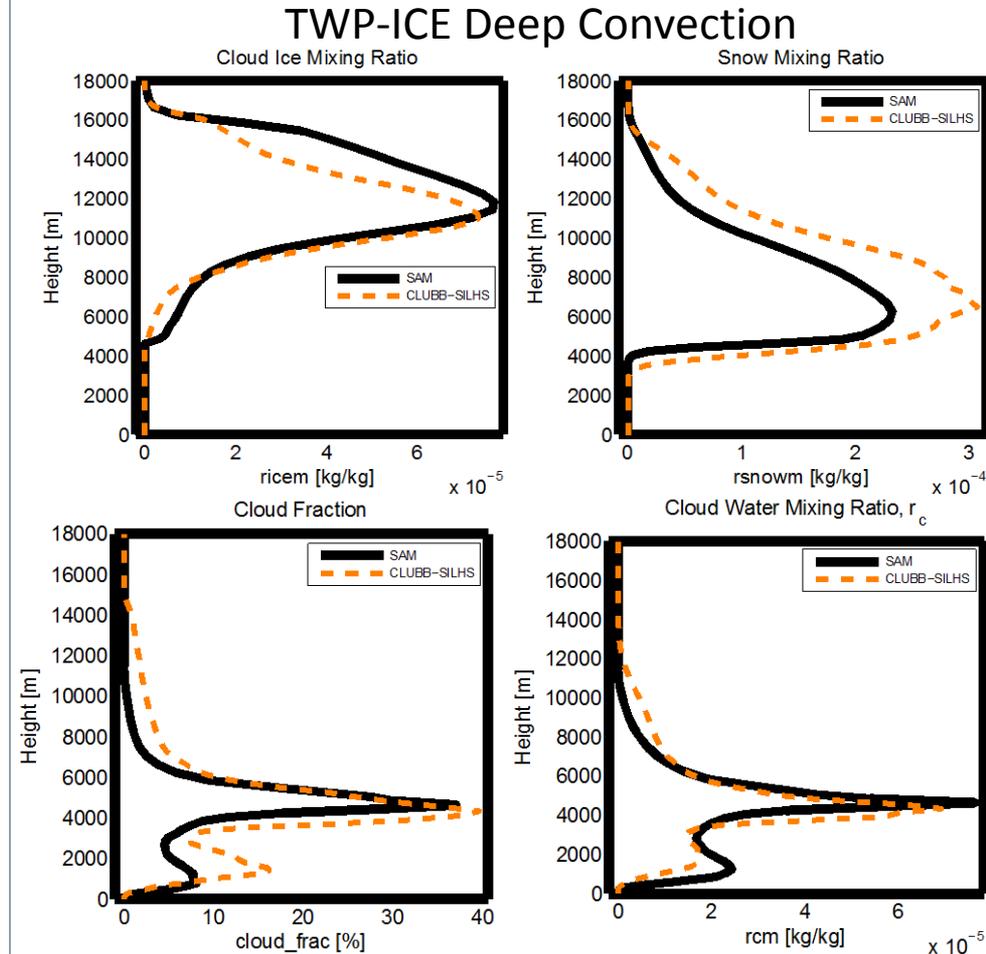
- Comparison of the residual-mean solution (TWA - top) to the bolus solution (GM - middle) is underway.
- An idealized configuration of the Antarctic Circumpolar Current is being used to compare/contrast TWA/GM.
- Equilibrated TWA system exhibits a meridional temperature gradient  $\sim 10\%$  stronger than the GM system (bottom).
- Tracking down root cause of differences. Likely due to different treatment of potential vorticity.



Mesoscale Eddy  
Treatments

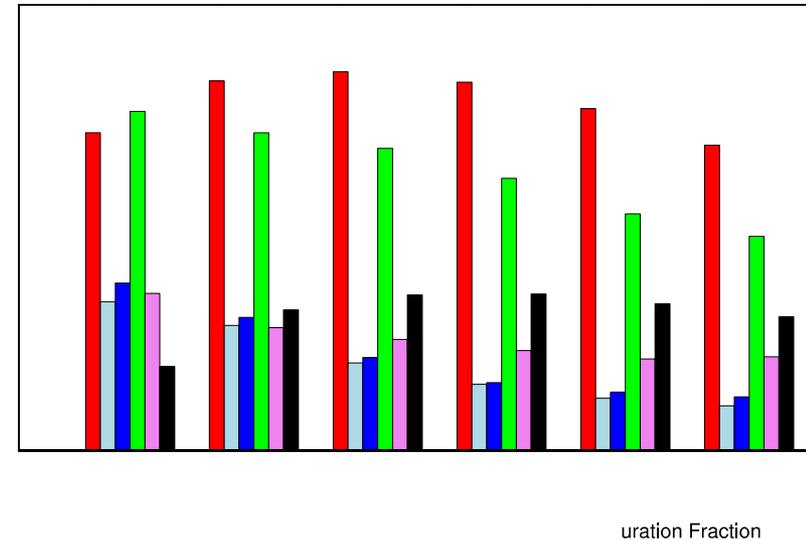
## Convection

- CLUBB cloud-turbulence probability density function scheme was designed for shallow clouds
- With two log-normal modes comprising the PDFs, it can represent skewness
- Using Latin hypercube sampling for microphysics, it simulates five deep, shallow and stratiform cloud cases realistically



## Convection

- Used 512 CAM5\_CLUBB simulations to determine the sensitivity of low cloud distribution to 18 tunable parameters
- Sensitivity analysis of effects of uncertainty in 11 CAM5 parameters on global mean RESTOM at 3 resolutions finds RESTOM uncertainty dominated by uncertainty in one parameter, with bias largest but variance smallest at finest resolution



- Evaluated six convection closures using cloud-resolving model output. Found that moisture convergence closure performs the best as model resolution increases

# Development of scale-aware cumulus scheme

Yi-Chin Liu<sup>1</sup>, Jiwen Fan<sup>1</sup>, Guang J. Zhang<sup>2</sup>, Kuan-Man Xu<sup>3</sup>, and Steve Ghan<sup>1</sup>

<sup>1</sup> Pacific Northwest National Laboratory; <sup>2</sup> Scripps Institution of Oceanography; <sup>3</sup> NASA Langley Research Center

## Objectives

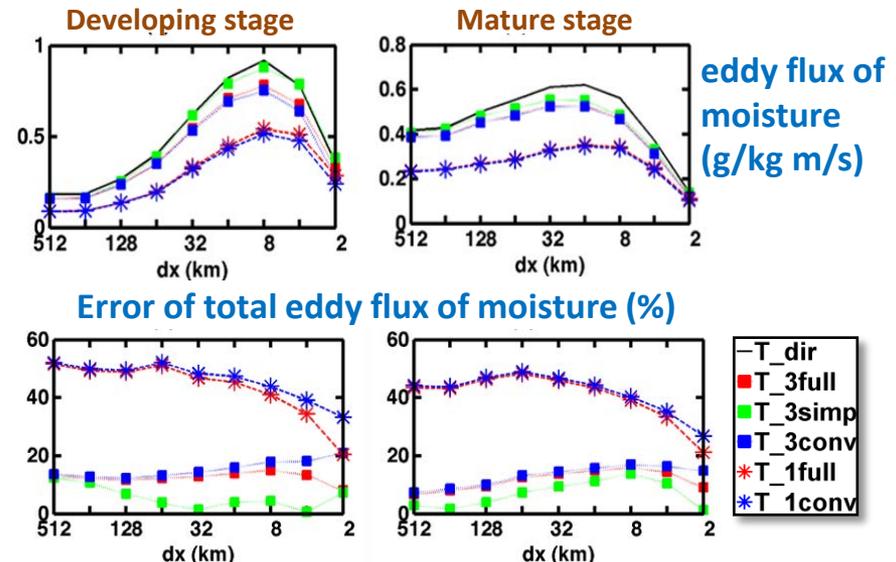
Develop a scale-aware cumulus parameterization based on the traditional Zhang-McFarlane (Z-M) scheme based on cloud resolving model simulations.

## Impact

More accurate clouds and water vapor in the upper troposphere at a wide range of model resolutions from mesoscale to global scale .

## Accomplishments

- Examine the scale-dependence of eddy transport of moisture at different altitudes, cloud life stages, and different regions.
- Evaluate different eddy transport formulations and propose a simplified formulation with 3-updraft and 1-downdraft, which has the following three advantages:
  - ✓ valid for cloud fraction up to 1
  - ✓ simple formulation
  - ✓ accurate representation of CRM-simulated eddy flux across scales



- Our proposed simplified 3-updraft formulation (**T\_3simp**), predicts eddy flux very close to CRM results (**T\_dir**). Single updraft approaches (**T\_1full**, **T\_1conv**) greatly underestimate eddy flux.

# Eddy flux of moisture in convection

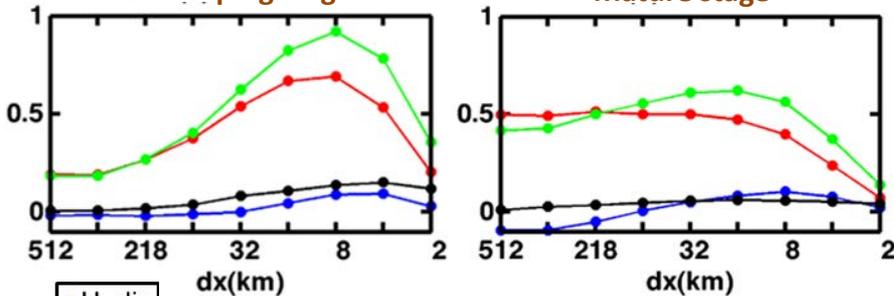
## Vertical eddy flux of moisture (g/kg m/s)

### Mid-latitude (MC3E)

5.5 km altitude

Developing stage

Mature stage

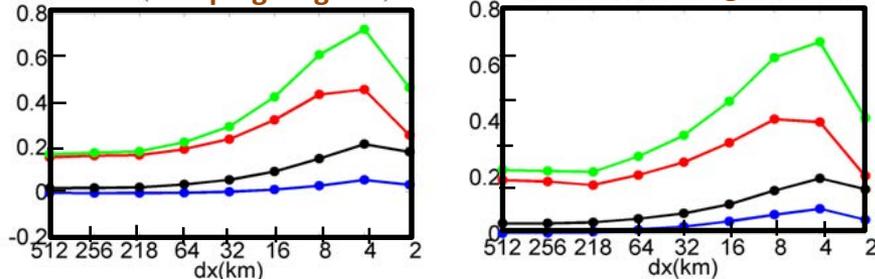


### Tropical region (TWP-ICE)

6.5 km altitude

Developing stage

Mature stage



Note: **U, D, E**: updraft, downdraft, and environment component  
**dir**: the direct calculation based on the CRM simulations.  
The **green line** is the total eddy flux

- Updraft eddy flux is the major contributor to total eddy flux
- Mid-latitude continental vs. tropical convection

### Mid-latitude convection:

- (a) Downdraft eddy transport can be as large as updraft eddy transport below 3-km altitude, especially at the mature stage.
- (b) Updraft eddy transport has stronger scale dependency at the developing stage than the mature stage.

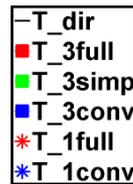
### Tropical convection:

- (a) Downdrafts contribute less to the total transport compared with the mid-latitude cases.
- (b) No distinguishable difference in scale dependency between the developing and mature stages.

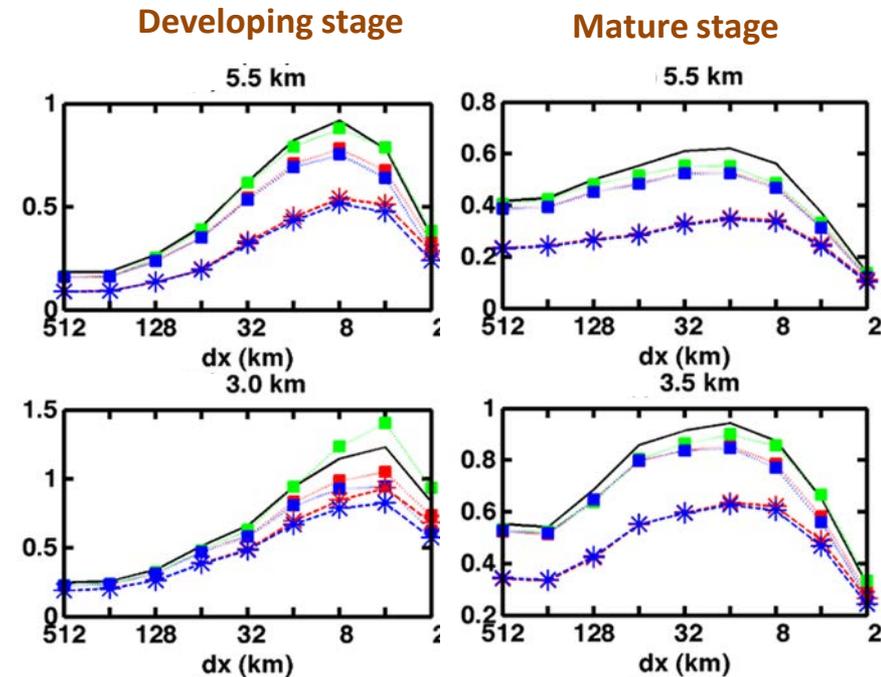
# New scale-aware eddy parameterization

We propose the simplified 3-updraft and 1-downdraft formulation for total eddy flux of moisture:

$$T_{3simp} = \sum_{k=1}^3 \sigma_{Uk} (1 - \sigma_{Uk}) (\widehat{w}_{Uk} - \widehat{w}_E) (\widehat{q}_{Uk} - \widehat{q}_E) + \sigma_D (1 - \sigma_D) (\widehat{w}_D - \widehat{w}_E) (\widehat{q}_D - \widehat{q}_E)$$



- The proposed formulation (green line) accounts for the inter-draft variance of updrafts and provides the best approximation to the direct CRM results (black line).
- Considering a cloud fractional area as proposed by Arakawa ( $T_{1full}$ ) does not significantly improve the eddy transport of moisture at any grid-spacing. Accounting for the inter-draft variability of updrafts is the key, especially at the scales of 4-64 km.



Y-axis: eddy flux of moisture (g/kg m/s)

Note:

$T_{dir}$ : total eddy flux from CRM simulations

**1 or 3**: 1-draft or 3-draft approach

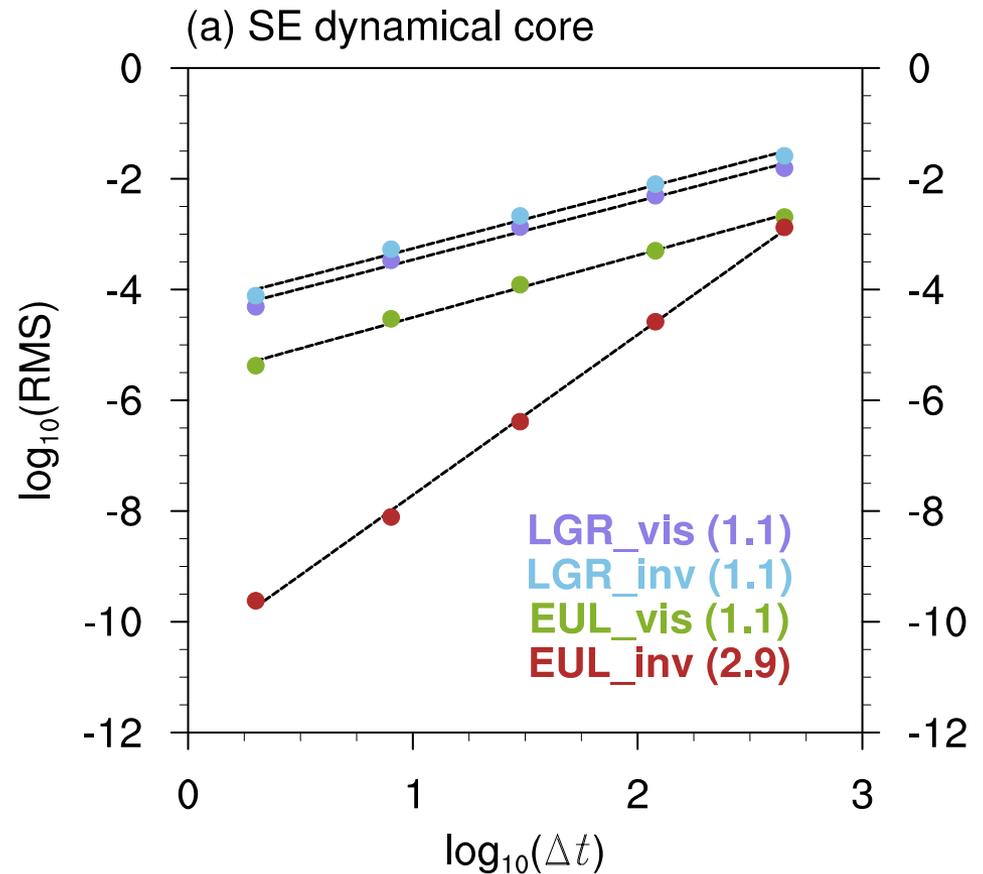
**conv**: formulation with assumption of cloud fraction  $\ll 1$

**full**: full formulation for a cloud fraction up to 1

**simp**: simplification of full neglecting interaction terms

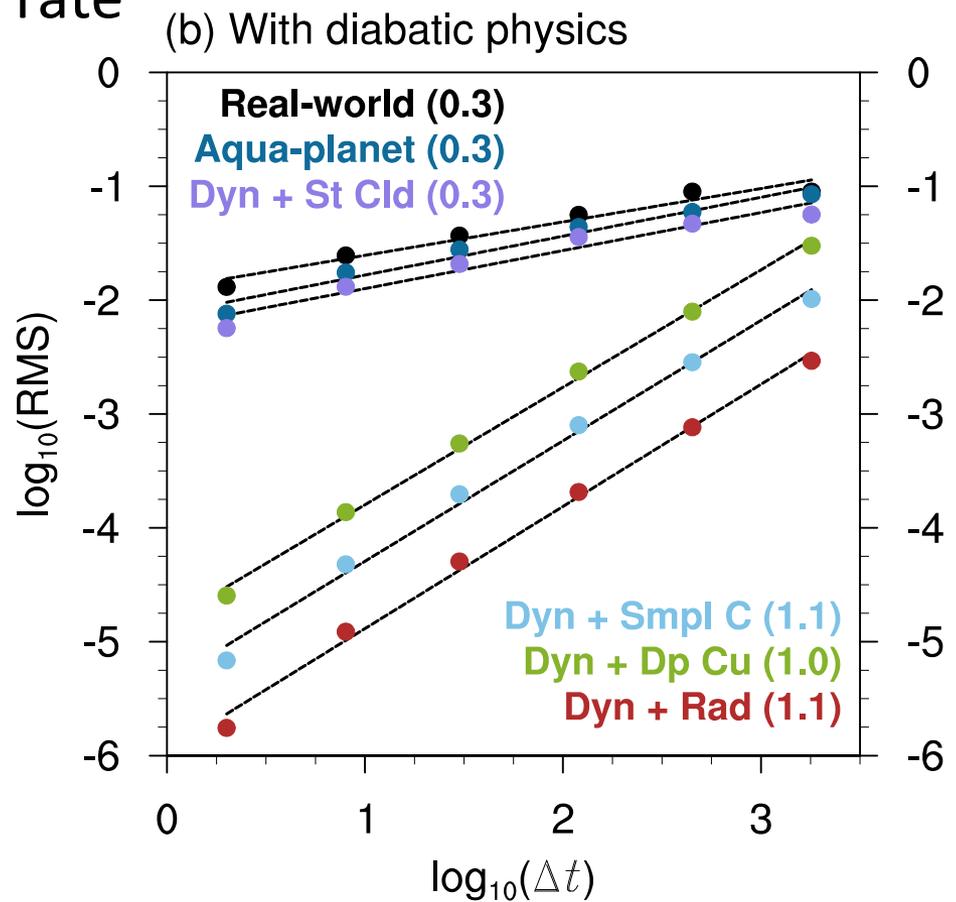
- Goal: Test physics convergence as  $\delta t \rightarrow 0$ :
  - Convergence rate (1<sup>st</sup> order as we expect?)
  - Magnitude of time stepping error
- Method: Evaluate physics convergence using:
  - Very short (1 h) simulations with CAM5.3
  - Time step sizes from 30 min down to 1 s
  - Metric: RMS difference of T

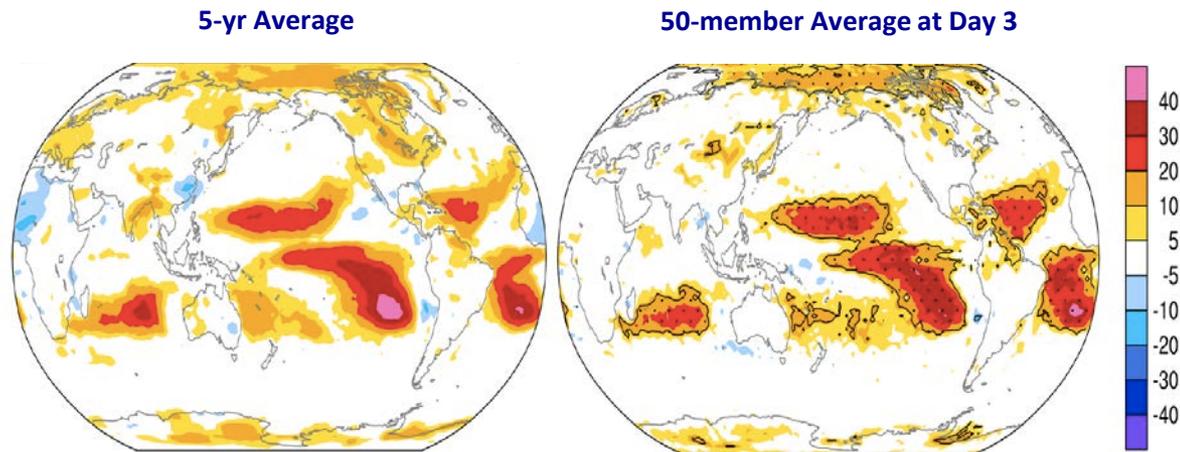
- Using only the dycore



# Impact of adding physics on convergence

- Full model converges much slower than expected
- Cloud macro and micro are most problematic
- In contrast, the simplified large-scale condensation converges at the expected rate





- Multiple simulation years are often required in sensitivity studies to overcome natural variability - inconveniently expensive at high resolutions
- We explored an alternative strategy using **ensembles of shorter simulations**, exploiting the important role of fast processes in determining model characteristics
- New method can correctly reproduce the main signals of model sensitivities revealed by long-term climate simulations, but at a fraction of total computation time and turnaround time.
- A powerful tool to efficiently use **flagship computing facilities** (e.g. Titan at Oak Ridge) and to **speed up model development**

## Sensitivity of Total Cloud Cover to Model Time Step (4 minute vs 30 minute)

Factor of 15 reduction in CPU time

*Reference:*

*Wan et al. (2014), Geosci. Model Dev. Discuss., 7, 2173-2216, doi:10.5194/gmdd-7-2173-2014*

# Multiscale Climate Application

- Progress due to SciDAC Institute interactions:
  - Implicit methods for time evolution at high resolution
  - Acceleration in ocean model performance
  - Addition of Chombo to suite of multiscale dycores
- Advances in the the physics / dynamics coupling:
  - Advances in the eddy formulation in our ocean model
  - Development of scale-aware convection schemes
  - Frameworks for ensuring physics time-step convergence